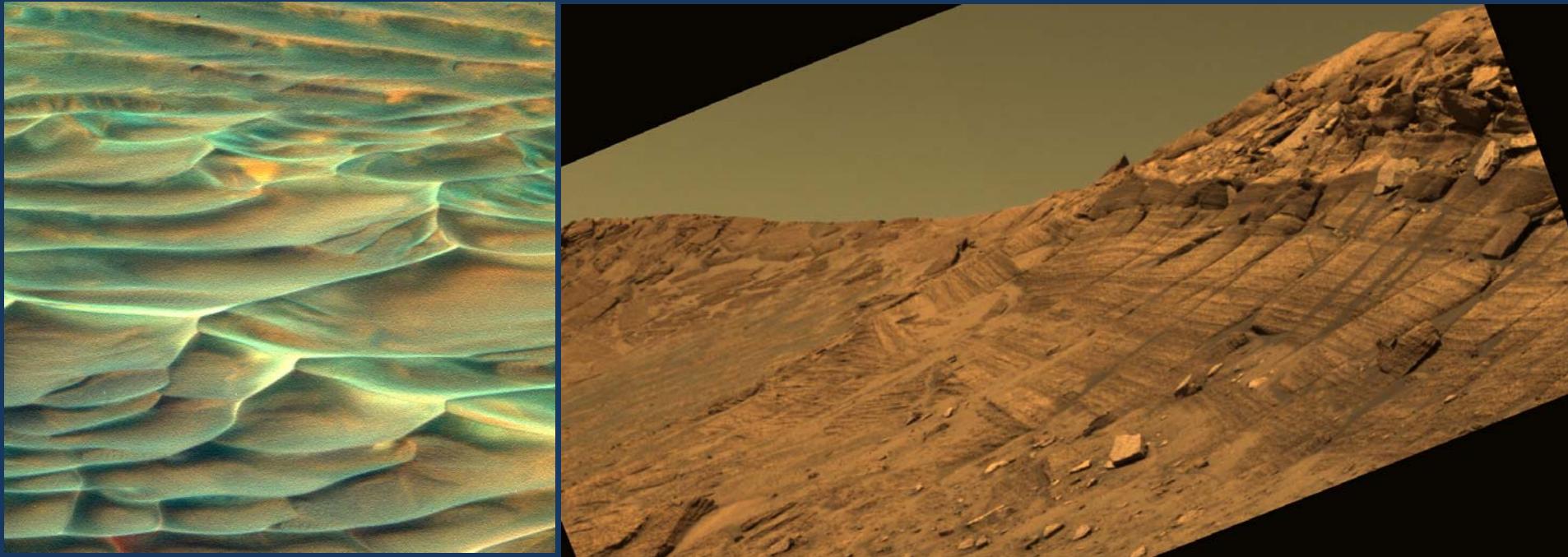


The Martian Sedimentary Mass: Constraints on its Composition, Age and Size

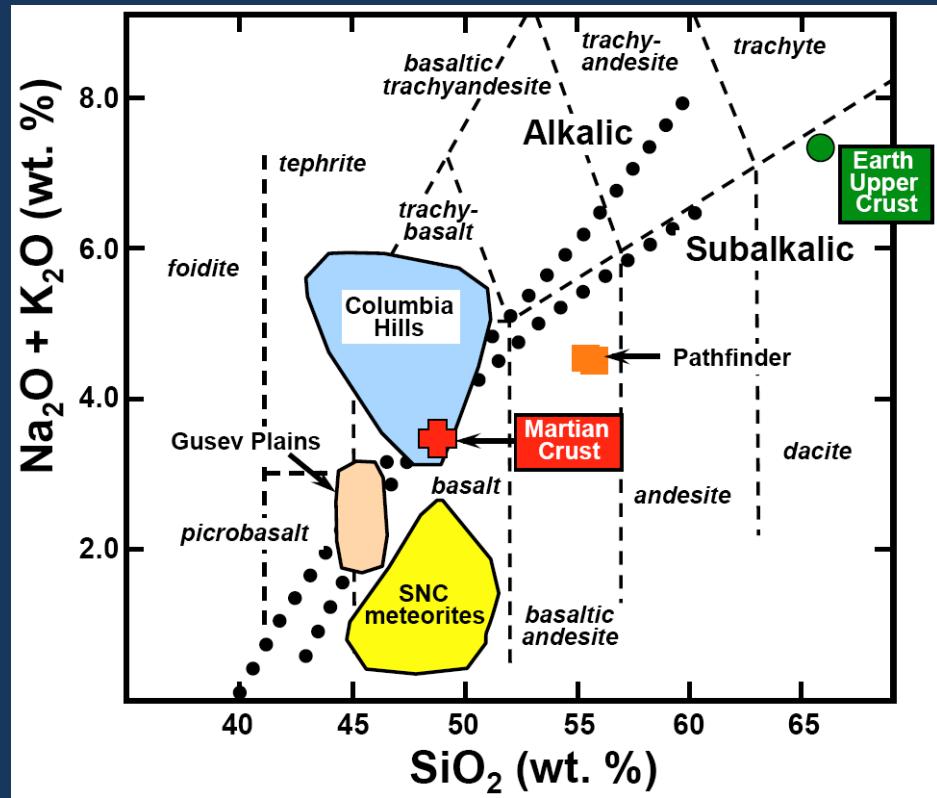
Scott McLennan

Department of Geosciences, SUNY Stony Brook



Martian Crustal Chemistry & Mineralogy

Taylor & McLennan (2009)



Mars Crust

17	Olivine	0
22	Pyroxene	1
10	Fe-Ti Oxides	1
22	Glass	13
29	Plagioclase	35
0	K-Feldspar	11
0	Sheet Silicates	14
0	Quartz	20

Earth Upper Crust

Olivine	0
Pyroxene	1
Fe-Ti Oxides	1
Glass	13
Plagioclase	35
K-Feldspar	11
Sheet Silicates	14
Quartz	20

McLennan & Grotzinger (2008)

SEDIMENTARY MINERALOGY

Earth

Mars

Phyllosilicates

illite (K)
smectite (Ca,Na,Mg)
kaolinite

CaCO_3
 $(\text{Ca},\text{Mg})\text{CO}_3$

CaSO_4

Fe_2O_3
 Fe_3O_4

Opal-A → Chert

montmorillonite (Mg)
nontronite (Fe,Mg)
kaolinite

CaCO_3
 MgCO_3

MgSO_4
 $\text{Fe}^{2+}\text{Fe}^{3+}_2(\text{SO}_4)_4$
 CaSO_4

Fe_2O_3

Opal-A

Carbonates

Sulfates

Iron Oxides

Silica

SEDIMENTARY MINERALOGY

Earth

Mars

Phyllosilicates

illite (K)
smectite (Ca,Na,Mg)
kaolinite

CaCO_3
 $(\text{Ca},\text{Mg})\text{CO}_3$

CaSO_4

Fe_2O_3
 Fe_3O_4

Opal-A → Quartz

montmorillonite (Mg)
vermiculite (Fe,Mg)

CaCO_3
 MgCO_3

$\text{Fe}^{3+}\text{Fe}^{2+}_2(\text{SO}_4)_4$
 CaSO_4

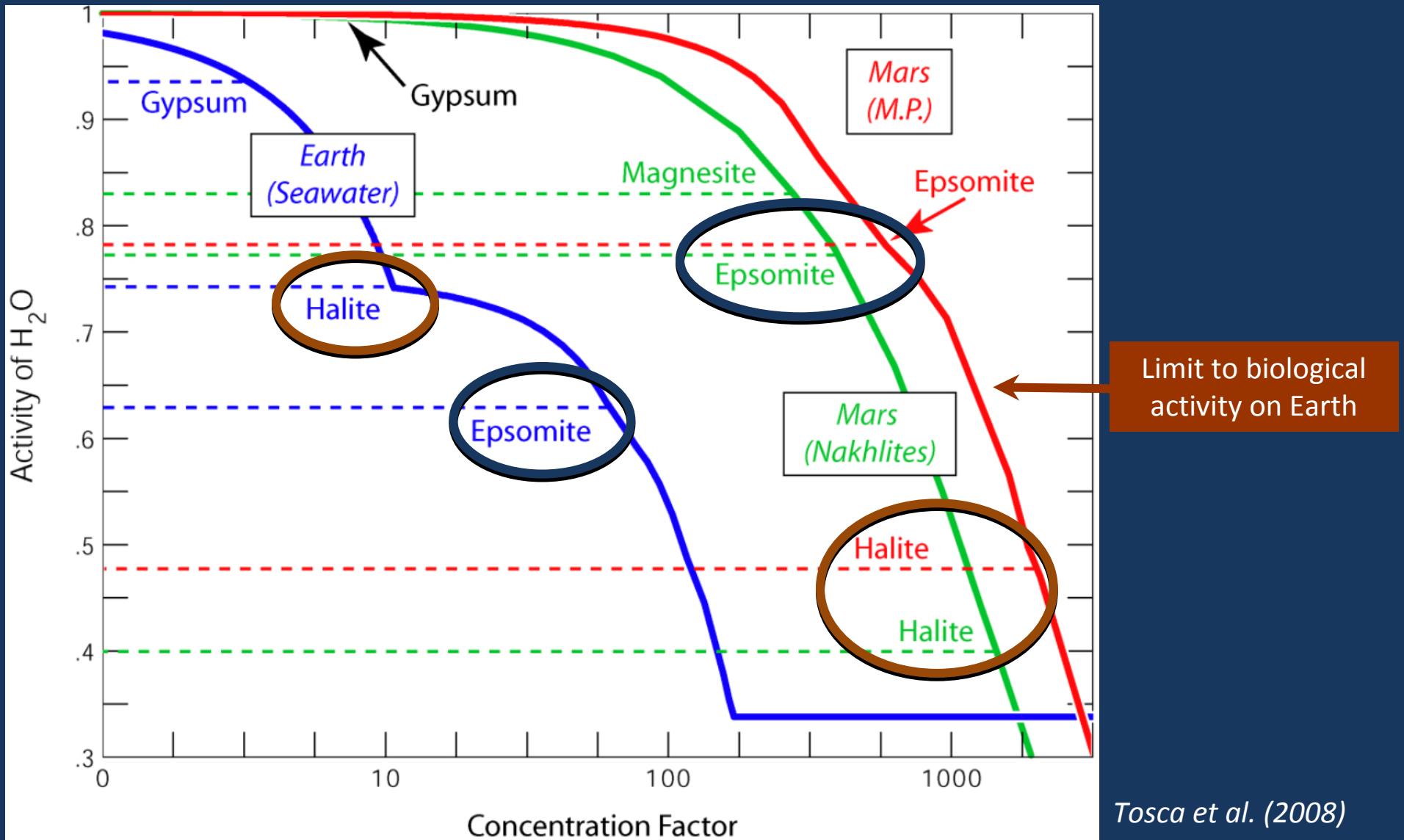
Fe_2O_3

Opal-A

Martian sedimentary minerals Fe-Mg-rich &
Na-K-poor compared to Earth reflecting
basaltic provenance;
“Immature” diagenetic assemblages

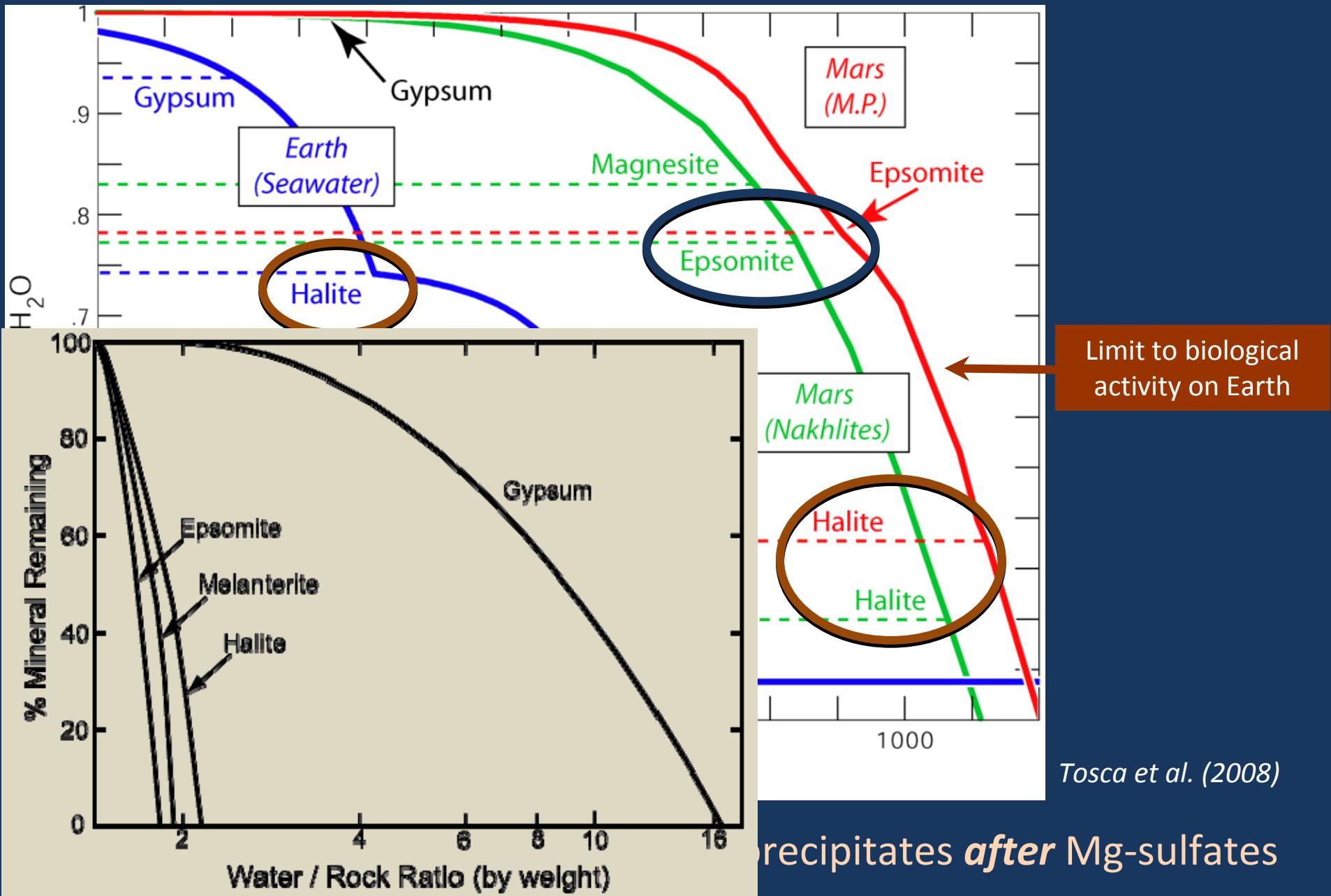
Silica

Evaporation Sequence

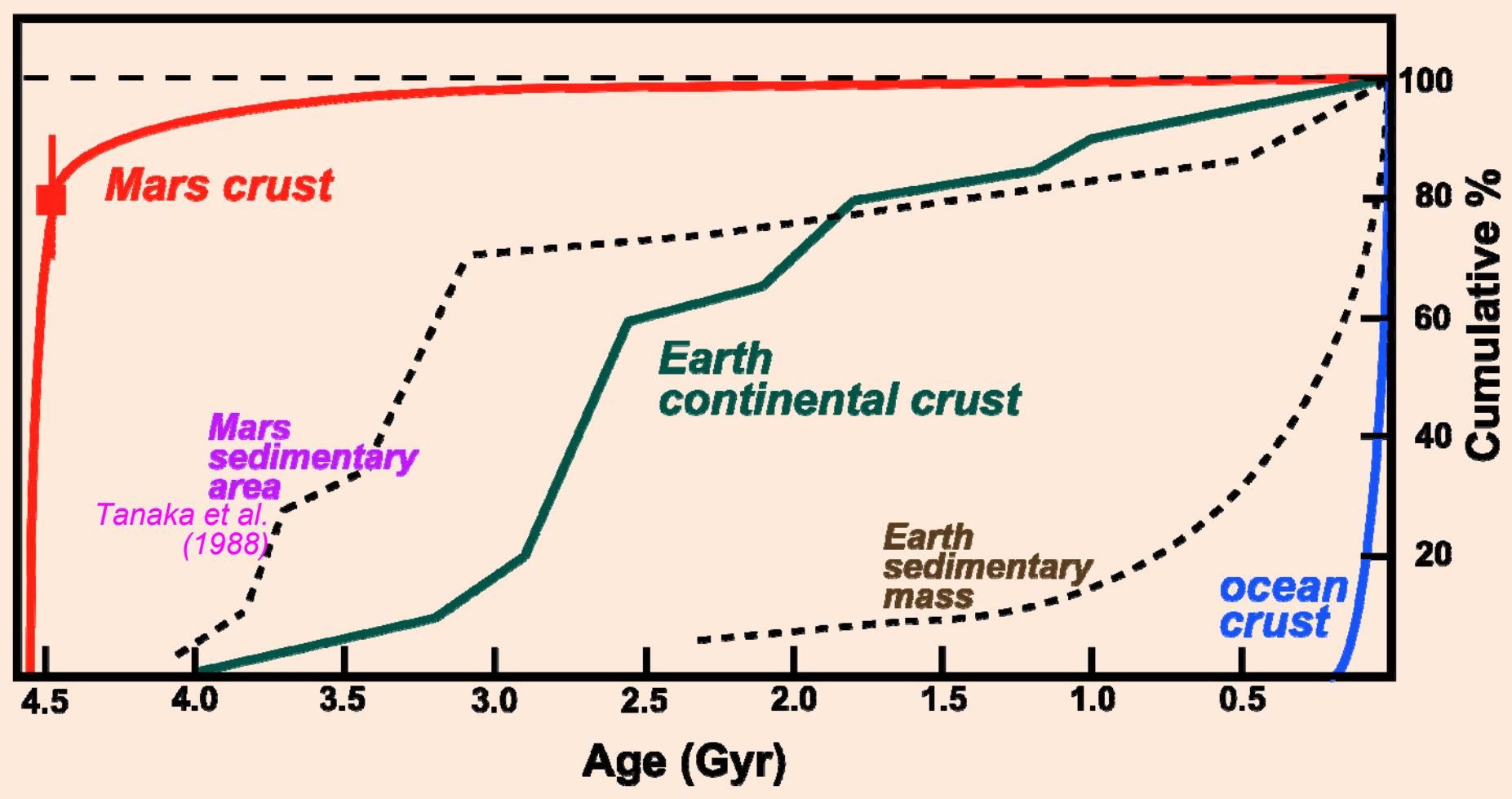


In S-rich Martian basaltic brines halite precipitates *after* Mg-sulfates

Evaporation Sequence



Crustal & Sedimentary Evolution

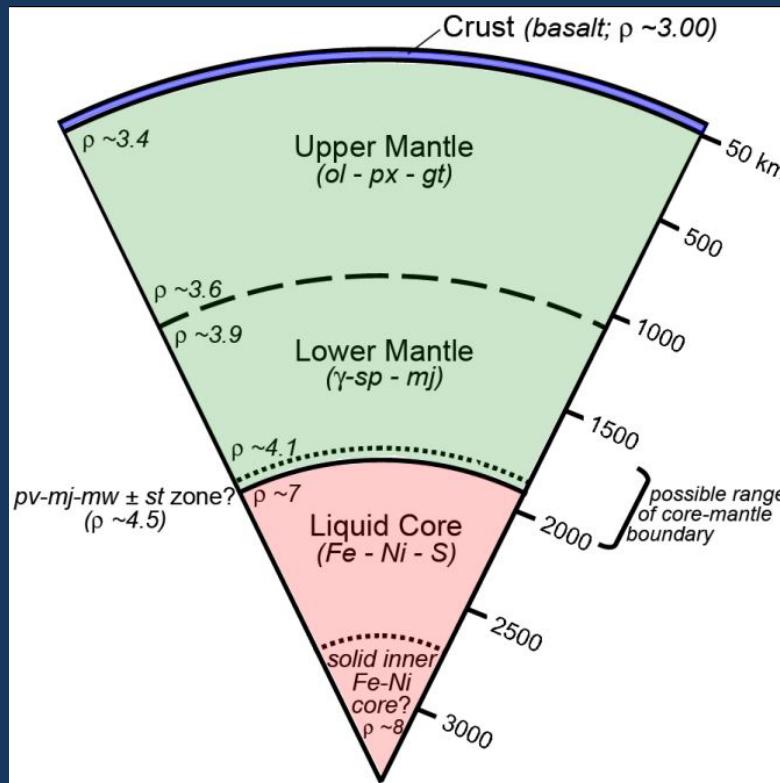


adapted from
Taylor & McLennan (2009)

How Large is the Martian Sedimentary Record?

- for comparison, terrestrial sedimentary record is $2.7 \times 10^{24} \text{ g}$
- direct measure from Martian stratigraphic records (as done on Earth) not possible
- indirect estimate from sulfur degassing history

Mars is a S-rich planet with crustal/mantle sulfur $\sim 2 \times$ terrestrial levels



Crust

$S \sim 2,000 \text{ ppm}$
($4.3 \times 10^{22} \text{ g}$)

Primitive Mantle

$S > 400 \text{ ppm}$
($> 2.1 \times 10^{23} \text{ g}$)

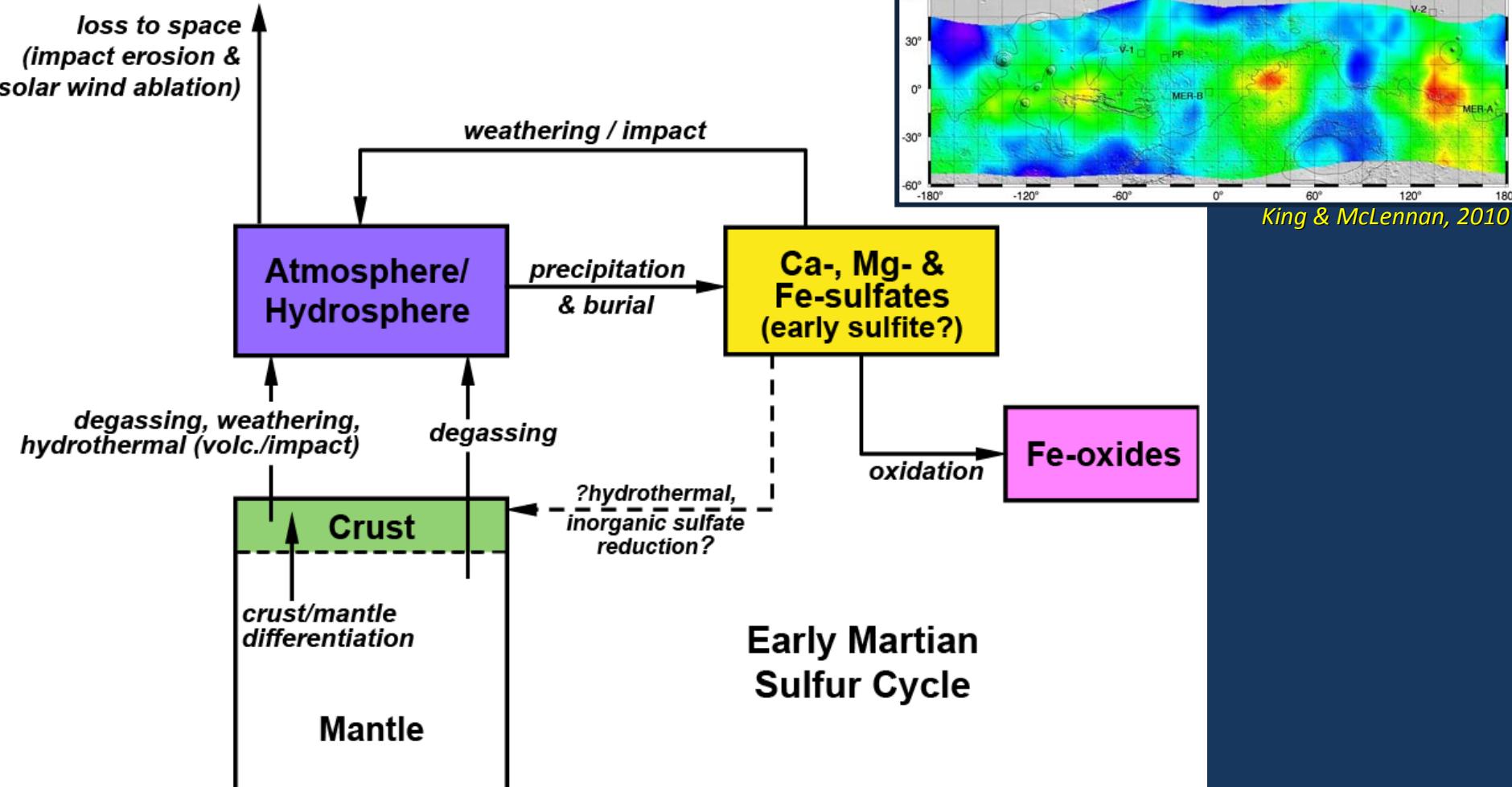
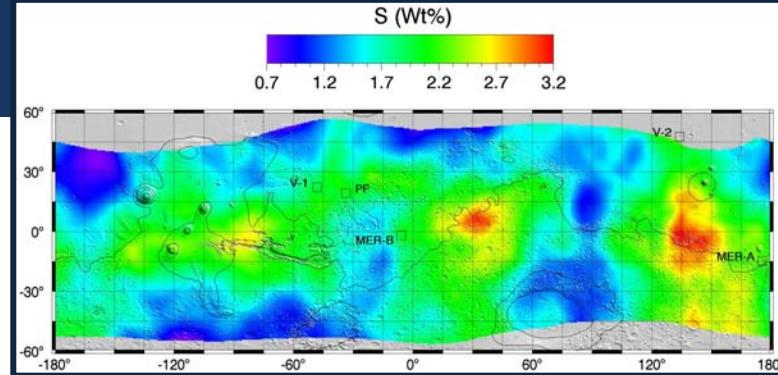
Core

$S \sim 14\%$

Ancient Sulfur Cycle on Mars

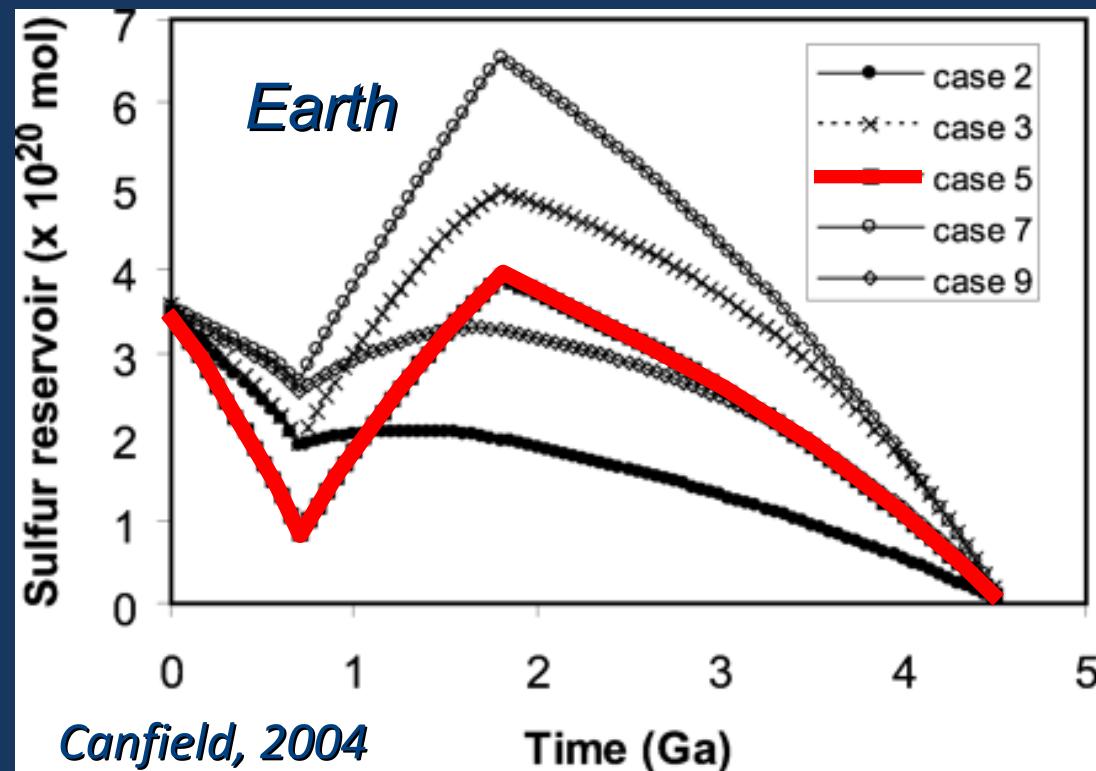
Mineral Epoch	Phyllosian	Theikian	Siderikian
Mineralogy	Clay-rich	Sulfate-rich	Anhydrous Ferric Oxide-rich
Environment	circum-neutral pH water-rich	low pH water-limited?	low pH water-limited
Age	Noachian	Hesperian	Amazonian
	~4.5Gyr	~3.7Gyr	~3.0Gyr
			0 Gyr

Bibring et al., 2006



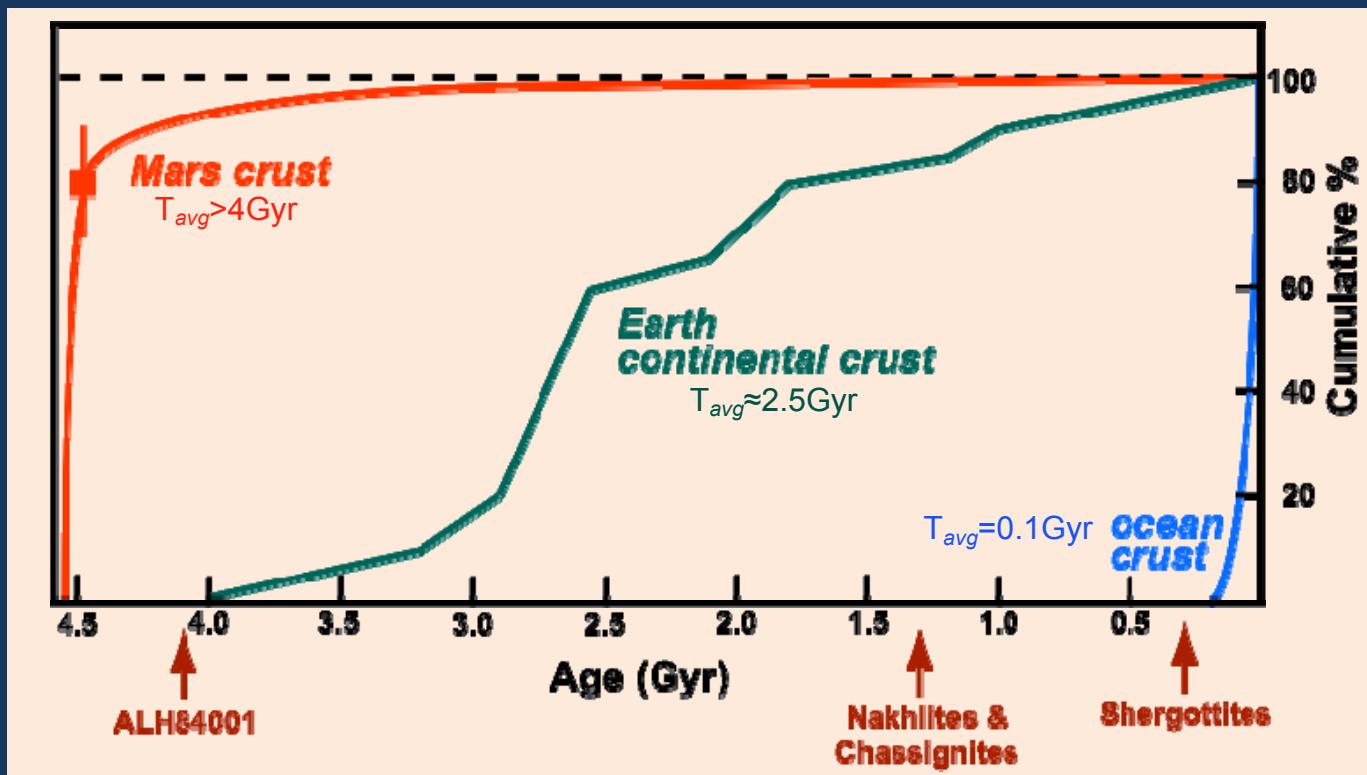
Sulfur Outgassing on Earth

- estimated size of Earth's outgassed S reservoir is model dependent
- Terrestrial history complicated by S loss from crust-mantle recycling (i.e., plate tectonics)
- estimate integrated total S through surface reservoir
- for modest assumptions
 - integrated over time
 $\sim 1.1 \times 10^{23}$ g
 - 11% of Earth's primitive mantle sulfur



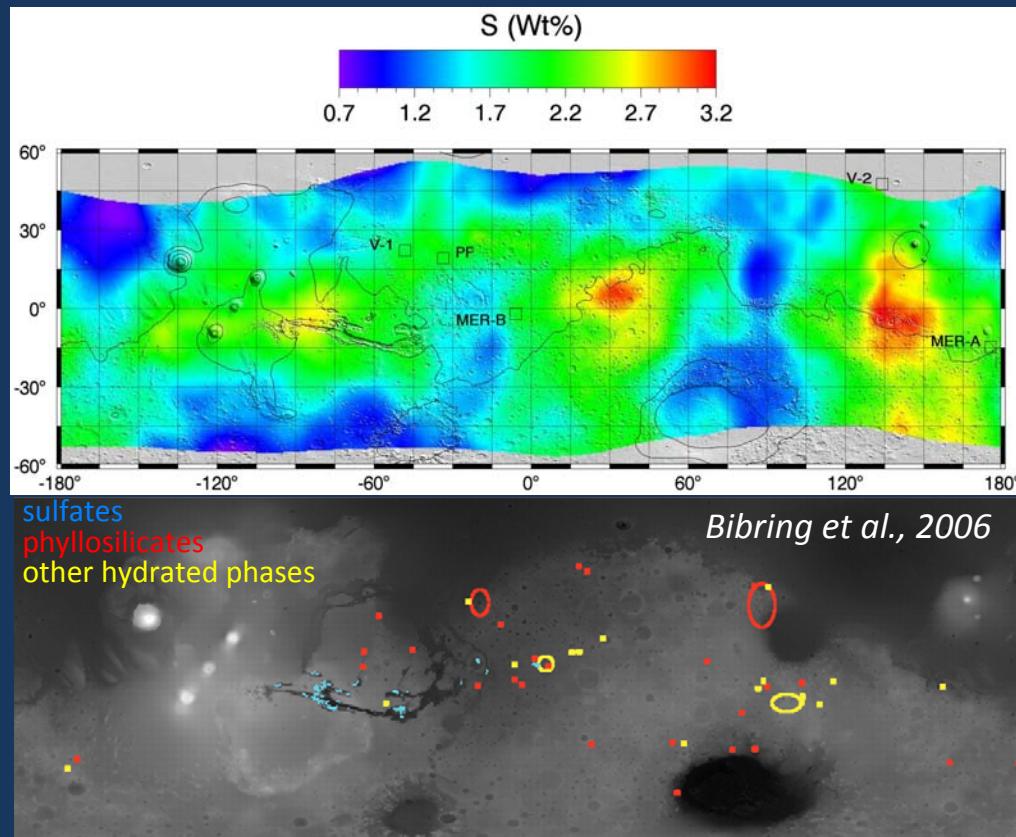
Sulfur Outgassing on Mars - 1

- Mars differentiated much earlier than Earth
- ≥50% differentiation of Mars primitive mantle versus ~25-30% for terrestrial mantle (and S is an incompatible element)
- no plate tectonics - no mantle recycling – what comes to the surface stays on the surface



Sulfur Outgassing on Mars -2

- assume 11% S outgassing – likely a lower limit
- adopt 400 ppm S in primitive mantle – also likely a lower limit
- Martian near-surface S reservoir:
 - $\sim 2.3 \times 10^{22}$ g from S outgassing
 - $\sim 2 \times$ current Earth value but $\sim 20\%$ integrated value
 - $\sim 2\text{-}3\text{km}$ with average GRS/soil composition
 - estimates from magmatic history $\sim 4\text{-}10 \times$ lower but all lower limits
 - $\geq 10^{21} - 10^{22}$ g Sulfur



Martian Sedimentary Mass -1

- For calculation, adopt lower S value of 10^{21} g
 - Chemical Sedimentary Mass:
 - For **sulfates** assume crustal cation proportions – $\text{Fe}_{0.4}\text{Mg}_{0.4}\text{Ca}_{0.2}\text{SO}_4 \cdot 2\text{H}_2\text{O}$
 - For **chlorides** assume soil S/Cl ratio (3.6) and 50:50 mix – $\text{NaCl} : \text{Mg}_{0.67}\text{Ca}_{0.33}\text{Cl}_2$
 - Assume negligible **carbonates**
 - For **sedimentary silica** assume one mole silica for each mole of sulfate and chloride – $\text{SiO}_2 \cdot 1.5\text{H}_2\text{O}$
- 5.5×10^{21} g

2×10^{20} g

2×10^{21} g

Olivine

$$(\text{Fe},\text{Mg})_2\text{SiO}_4 + 4\text{H}^+ = \text{Fe}^{2+} + \text{Mg}^{2+} + 2\text{H}_4\text{SiO}_4$$

Clinopyroxene

$$\text{FeMgSi}_2\text{O}_6 + 4\text{H}^+ + 2\text{H}_2\text{O} = \text{Fe}^{2+} + \text{Mg}^{2+} + 2\text{H}_4\text{SiO}_4$$

H_2SO_4
 HCl

$\text{Fe}_2\text{O}_3, \text{FeSO}_4, \text{etc.}$

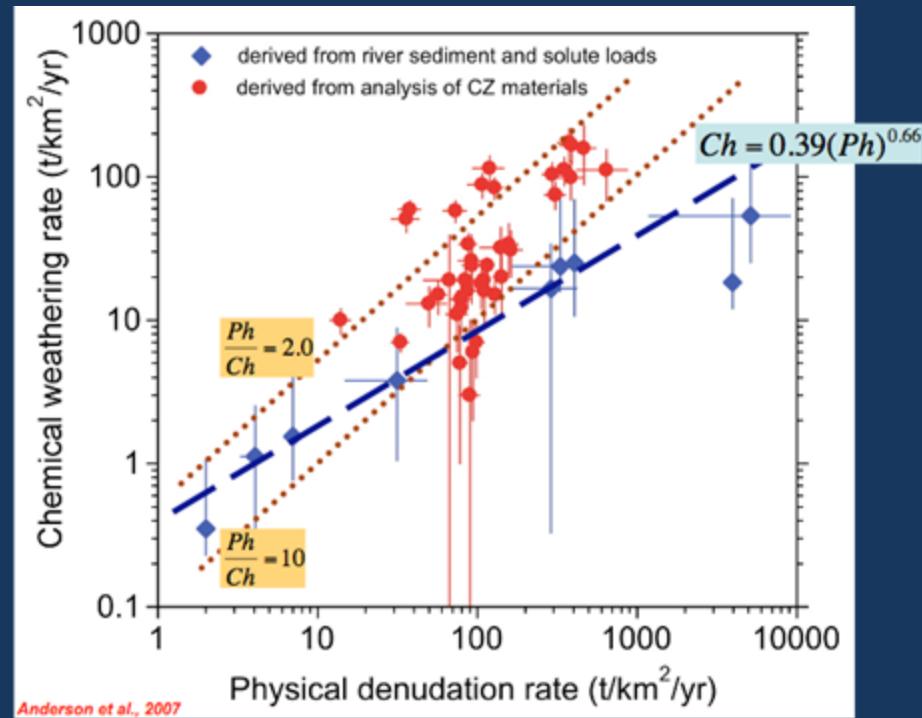
$\text{MgSO}_4, \text{MgCl}_2, \text{etc.}$

$\text{SiO}_2, \text{Siliceous Clays}^*$

8×10^{21} g
- Total Chemical Sedimentary Mass:

Martian Sedimentary Mass -2

- Ratio between clastic (Ph) & chemical (Ch) constituents studied on Earth – estimates for total sedimentary mass range from **3:1** to **6.5:1**
- Ratio likely higher on Mars – greater role for impacts and pyroclastics providing particulates
- Assume $Ph/Ch = 5$
- Indicates Martian sedimentary mass **5×10^{22} g** for 10^{21} g S and **5×10^{23} g** for 10^{22} g S
- **Martian sedimentary mass is $\geq 2\% - 20\%$ terrestrial sedimentary mass**



Some Conclusions

- Martian sedimentary record has distinctive chemical and mineralogical character compared to Earth due to basaltic crust:
 - *chemical and clastic minerals and sedimentary rock fragments are Fe-Mg-rich*
 - *chemical precipitation pathways differ (e.g., evaporative minerals and evaporation sequences)*
- Martian sedimentary mass much older than terrestrial sediments due to early formed crust and probable distinctive sedimentary recycling history
- Assuming degassed sulfur reacts to form chemical mineralogical constituents, it is possible to estimate the overall Martian sedimentary mass
 - $\geq 5 \times 10^{22} - 5 \times 10^{23}$ g
 - $\geq 2 - 20\%$ of the size of the terrestrial sedimentary mass