

Subsurface Aqueous Alteration on Ancient Mars: Implications for Habitability

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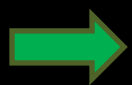
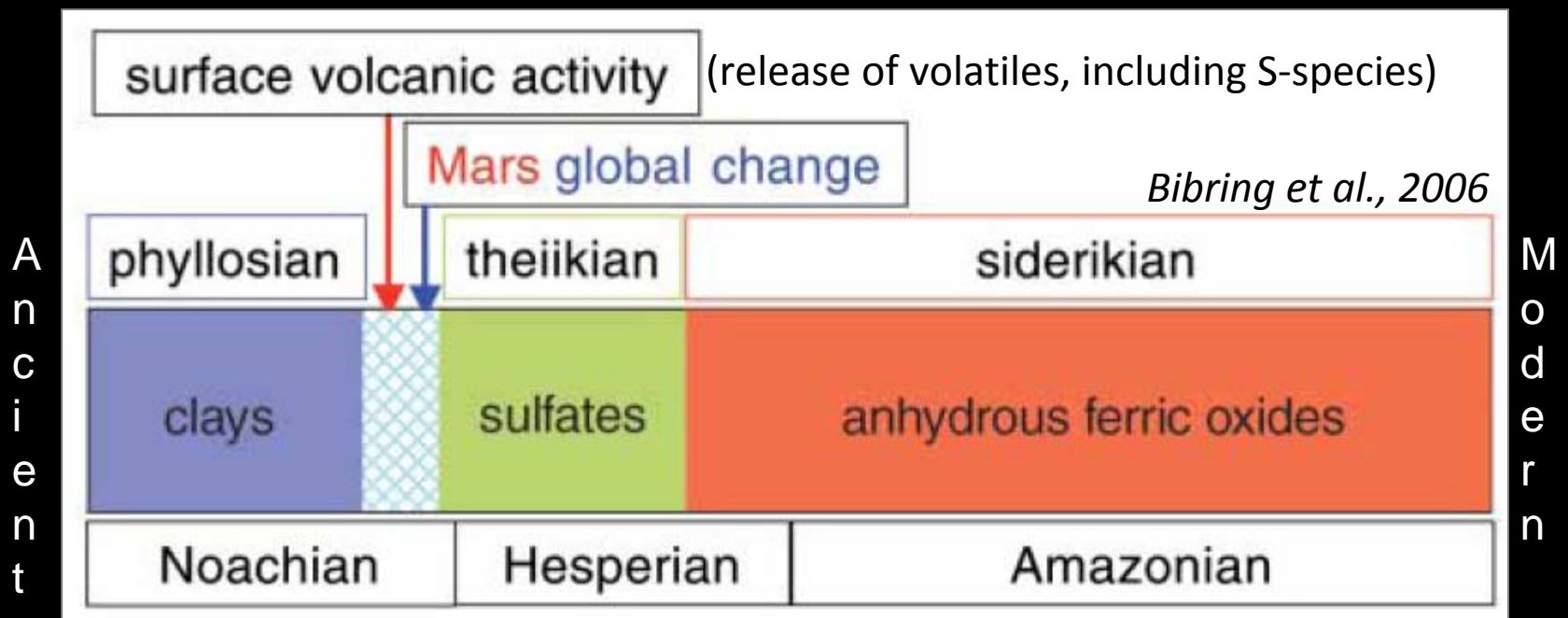
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“If phyllosilicates had formed in the subsurface, the Mars environment might have always been tenuous, cold, and dry, except for transient episodes...”

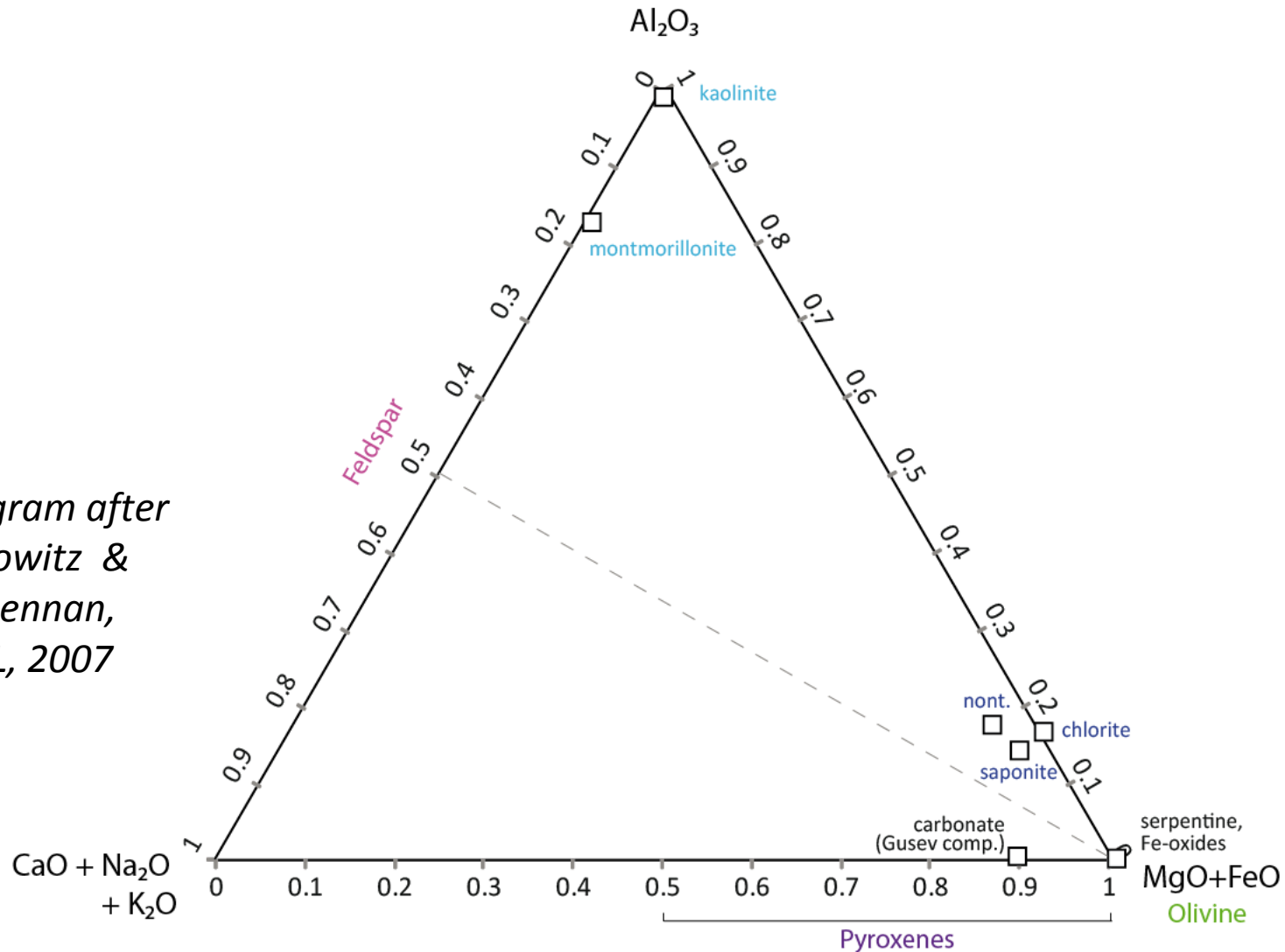
If instead **phyllosilicates formed at or close to the surface**, this would **require the Mars early atmosphere to be dense**. The global change to an acidic environment would then have been coupled to a rapid drop in atmospheric pressure.”

How to distinguish alteration environments?

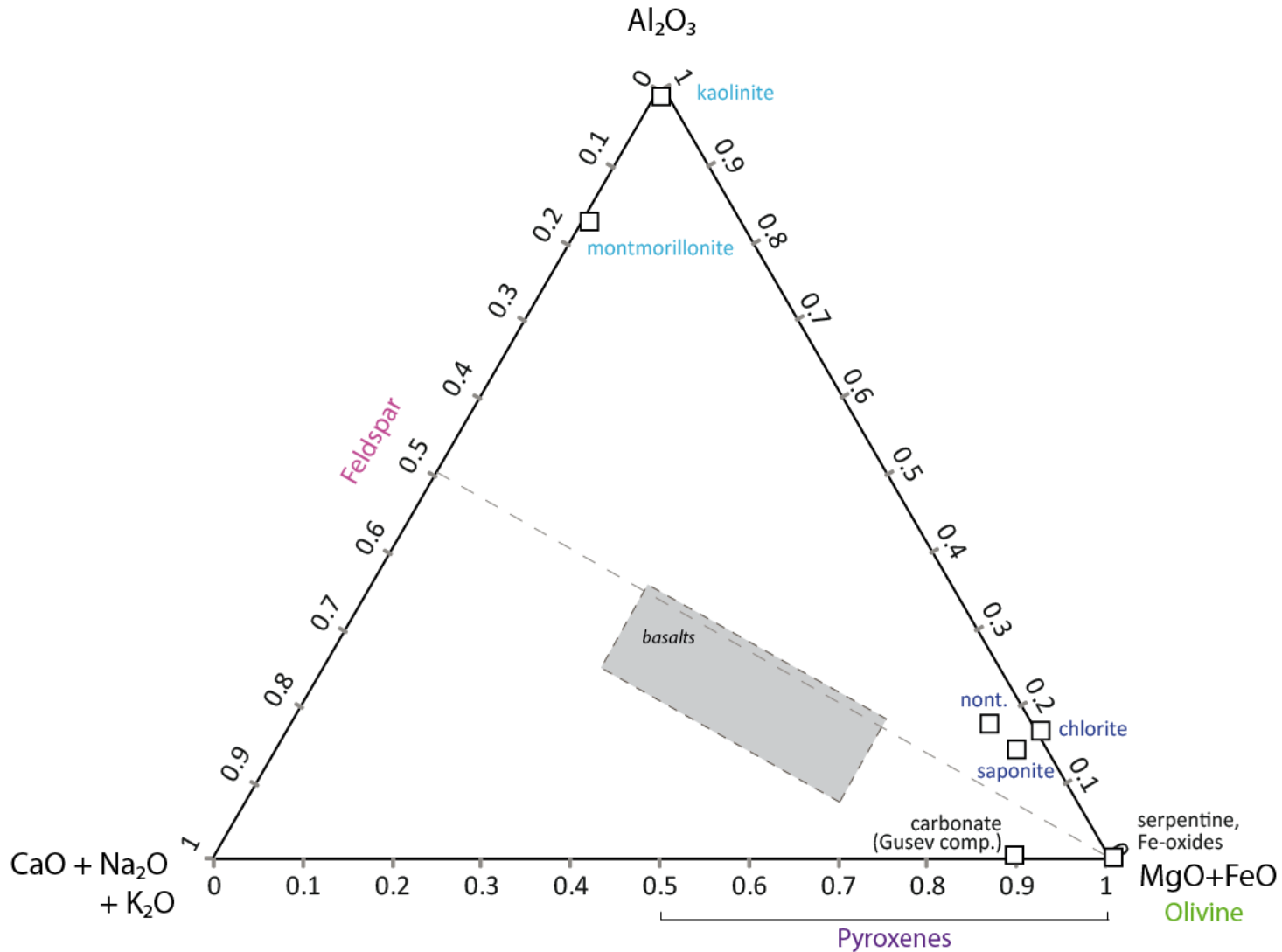
Geochemical and Mineralogic Indicators under
Chemically “Open” and “Closed” conditions
and low and high Water:Rock ratios

Styles of Alteration

*Diagram after
Hurowitz &
McLennan,
EPSL, 2007*



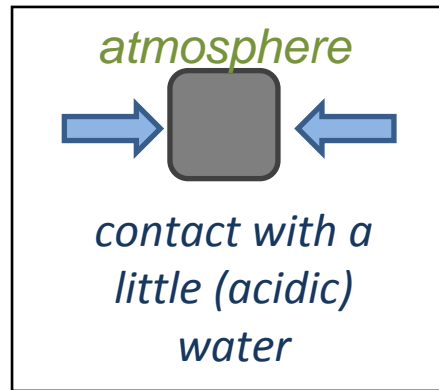
Styles of Alteration



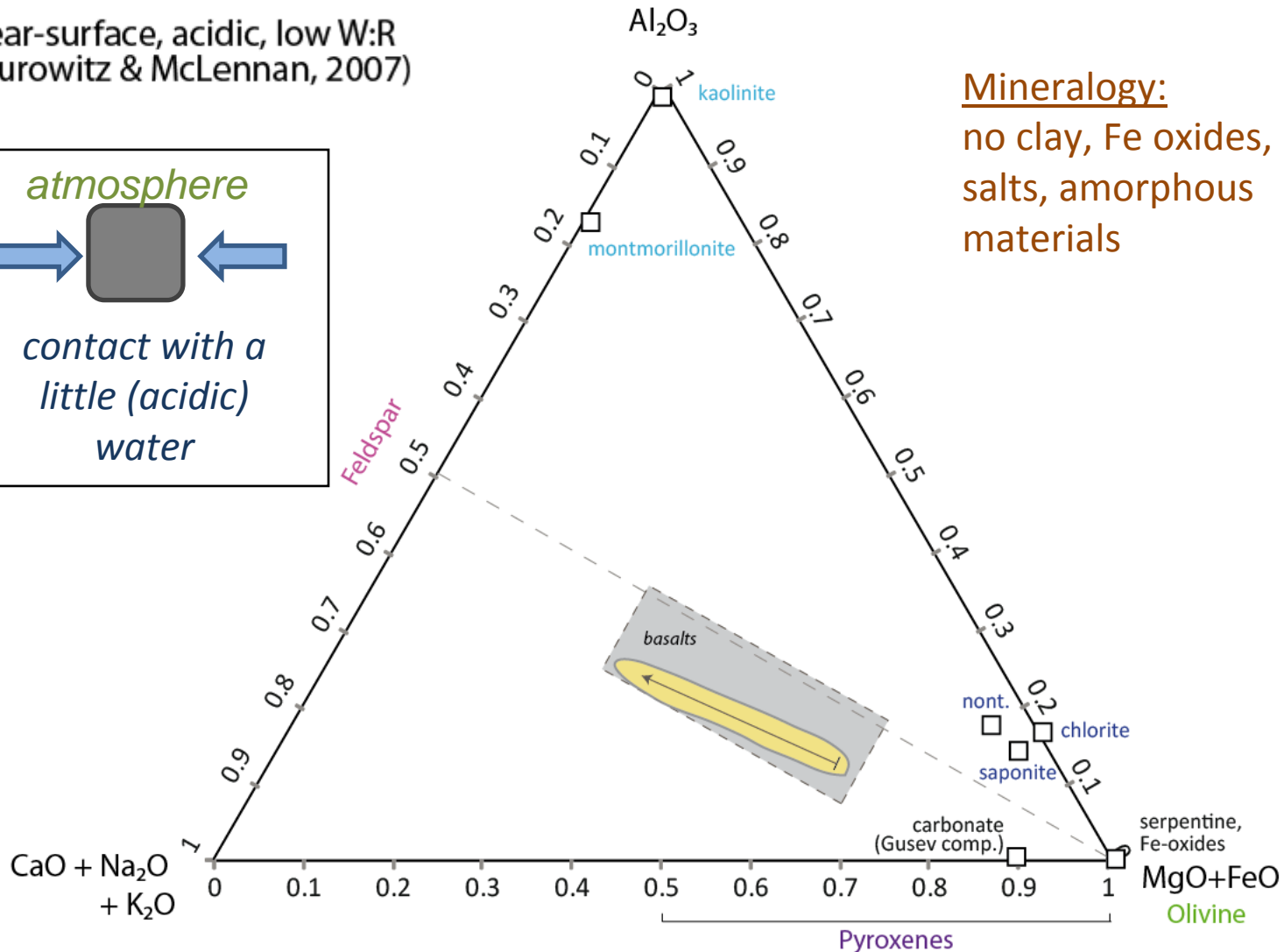
Styles of Alteration to form <not clay>



Near-surface, acidic, low W:R
(Hurowitz & McLennan, 2007)

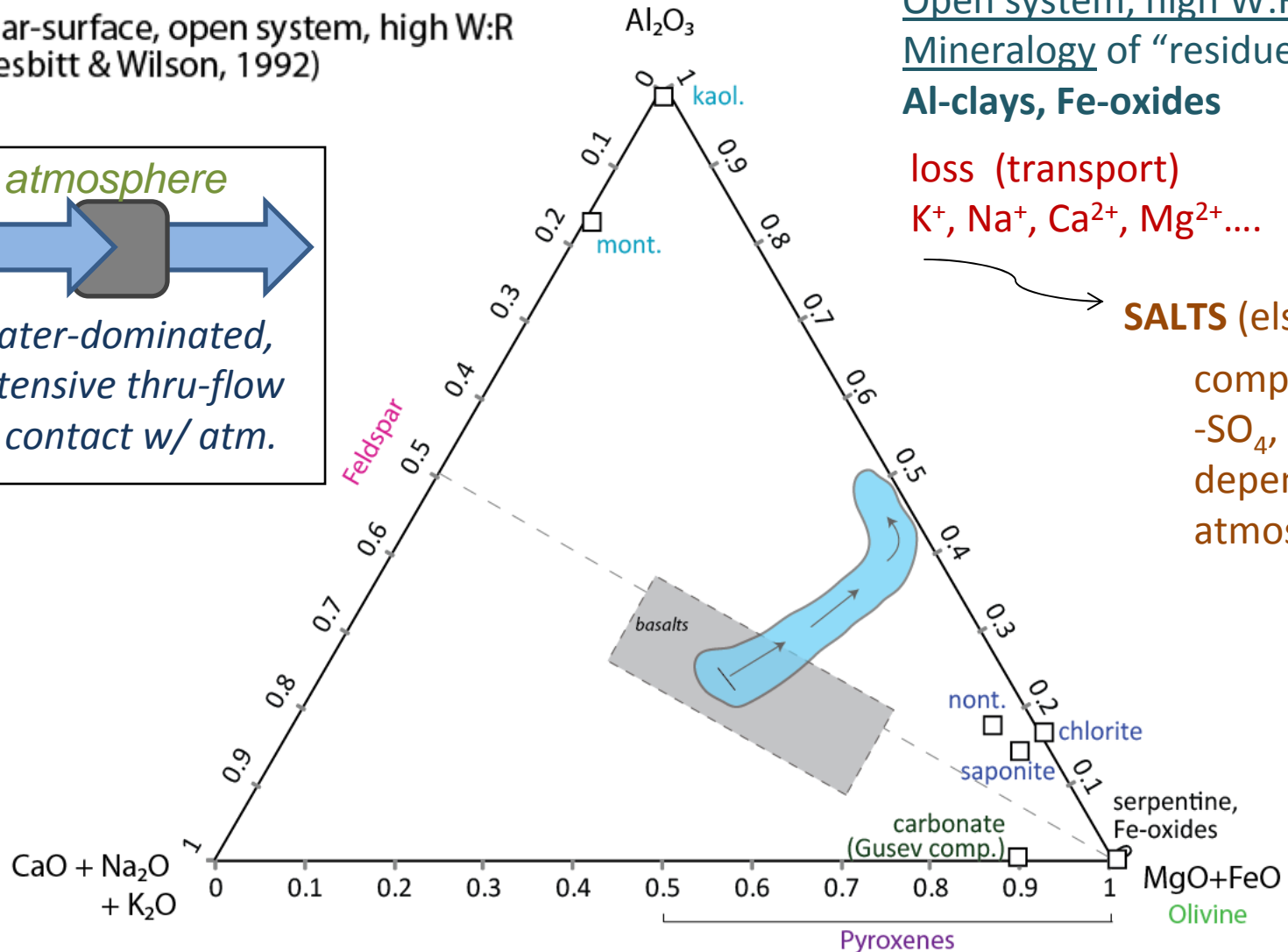
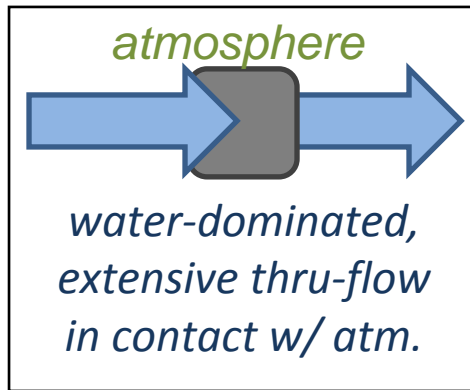


Mineralogy:
no clay, Fe oxides,
salts, amorphous
materials



Styles of Alteration to form Clays

- Near-surface, open system, high W:R (Nesbitt & Wilson, 1992)



Open system, high W:R
Mineralogy of "residue":

Al-clays, Fe-oxides

loss (transport)
K⁺, Na⁺, Ca²⁺, Mg²⁺....

SALTS (elsewhere)

composition...
-SO₄, -Cl, -CO₃,
depends on
atmosphere

Styles of Alteration to form Clays

- Open system, high W:R, near surface
(Nesbitt & Wilson, 1992)
- Subsurface, closed system
(hydroT alt. ocean ridge)
(Cann & Vine, 1966)
- Subsurface, closed system
(hydroT alt., terrestrial Iceland)
(Ehlmann et al., 2010)

Low W:R, closed system

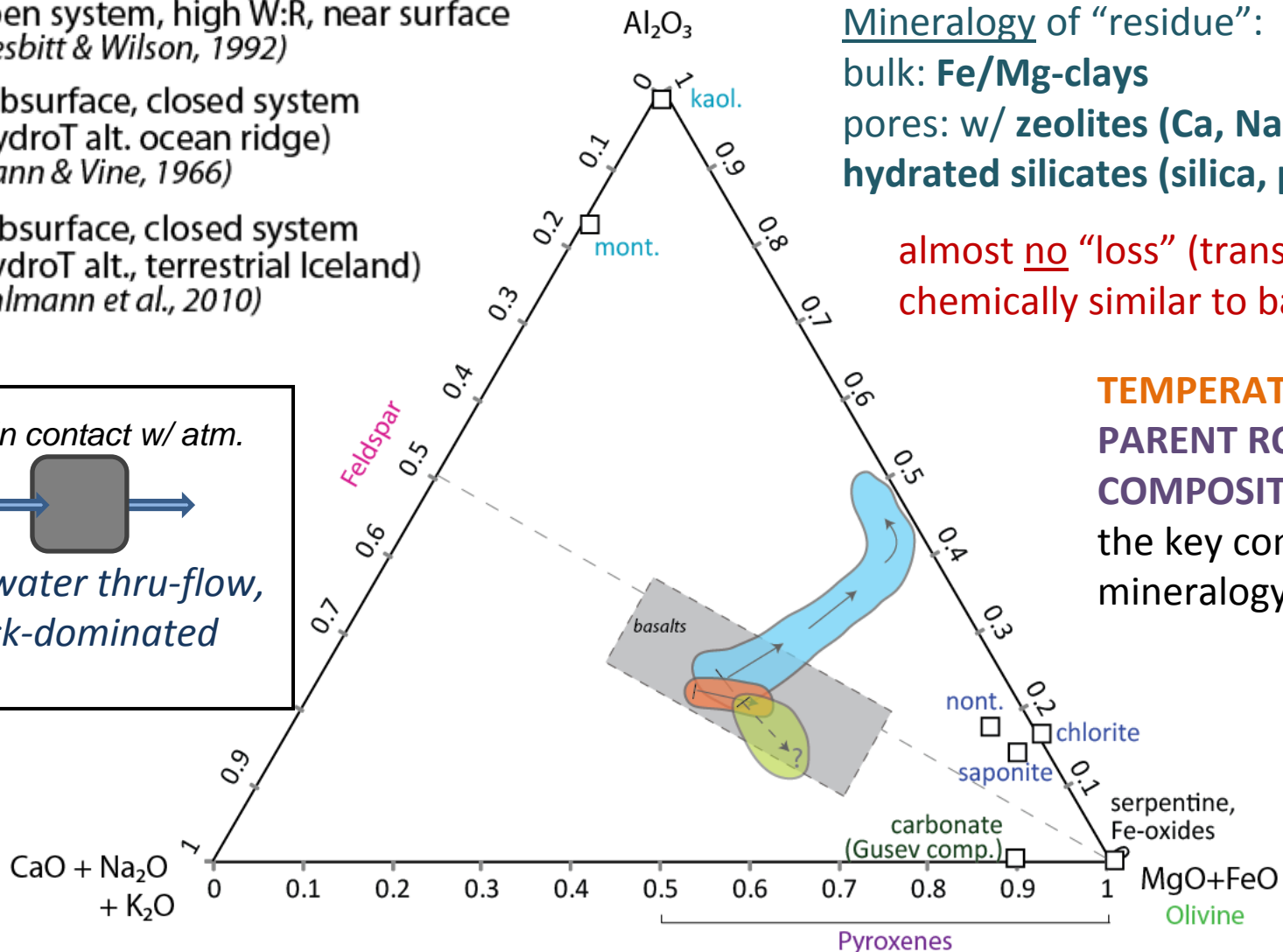
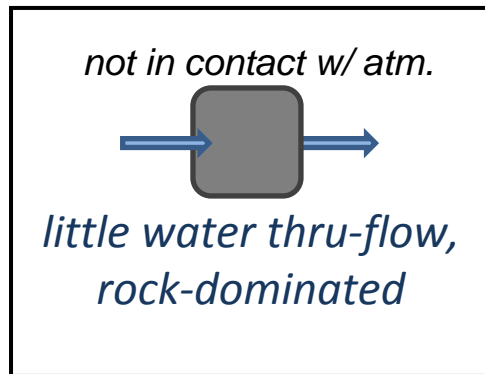
Mineralogy of “residue”:

bulk: **Fe/Mg-clays**

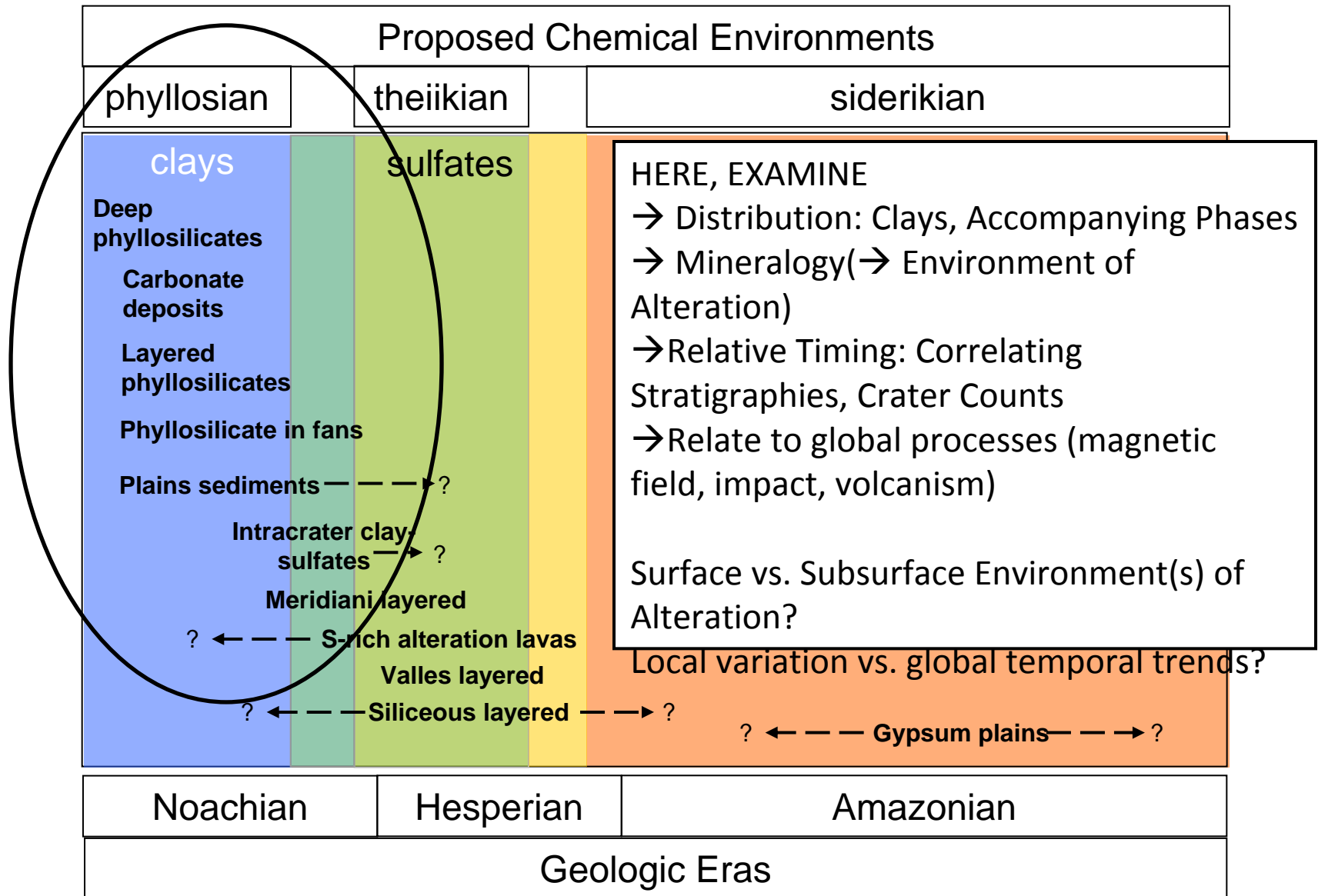
pores: w/ **zeolites (Ca, Na)**, other
hydrated silicates (silica, prehnite)

almost no “loss” (transport),
chemically similar to basalt

TEMPERATURE,
PARENT ROCK
COMPOSITION are
the key controls on
mineralogy



Adapted from Murchie et al., JGR, 2009 with mineralogic epochs from Bibring et al., 2006



A meta-analysis of CRISM data

DATA SOURCES

- Surveyed targeted images (18-40m/pixel)
- Compiled data from the literature when a complete list of alteration minerals found within a CRISM targeted image was reported by the authors

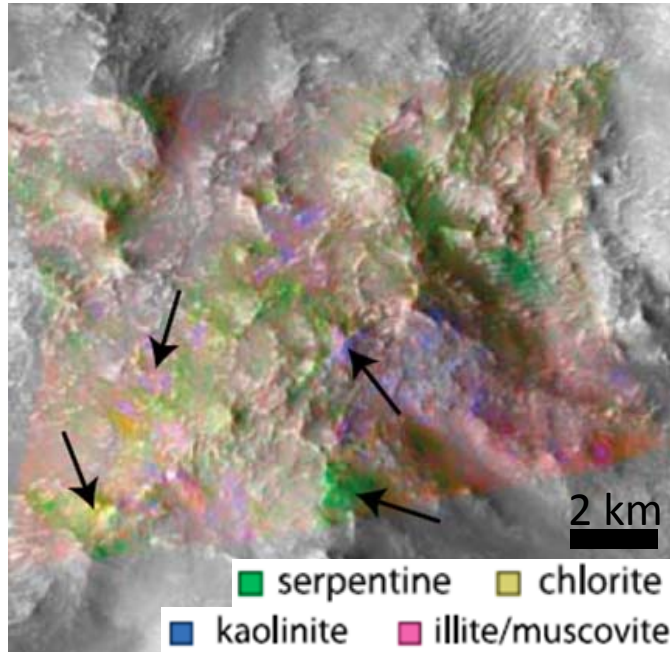
DATASET

- Lat/Lon of image center (images within +/- 15km count as 1 image)
- Kept track of present/absence of phyllosilicate and which accompanying phases
- Defined setting: “crustal”, “clay in stratigraphy”, “sedimentary”

Number of Data Points	Reference
1	Baldrige et al., 2009, <i>GRL</i>
6	Buczkowski et al., 2010, <i>JGR</i>
6	Buczkowski et al., 2010, <i>LPSC</i>
91	Carter et al., 2010, <i>Science</i>
75	Ehlmann et al., 2009, <i>JGR</i>
119	surveyed as part of Ehlmann et al., 2010, <i>GRL</i>
240	surveyed as part of Fraeman et al., 2009, <i>LPSC</i>
5	Glotch et al., 2010, <i>GRL</i>
1	Michalski & Niles, 2010, <i>Nature Geosci.</i>
26	McKeown et al., 2009, <i>JGR</i>
2	Milliken et al., 2010, <i>GRL</i>
4	Milliken and Bish, 2010, <i>Phil. Trans.</i>
2	Murchie et al., 2009, <i>JGR</i>
1	Mustard & Ehlmann, 2011, <i>LPSC</i>
31	Noe Dobrea et al., 2010, <i>JGR</i>
2	Roach et al., 2010, <i>Icarus</i>
5	Wiseman et al., 2010, <i>JGR</i>
1	Wiseman et al., 2008, <i>GRL</i>
21	Wray et al., 2011, <i>JGR</i>
639	TOTAL

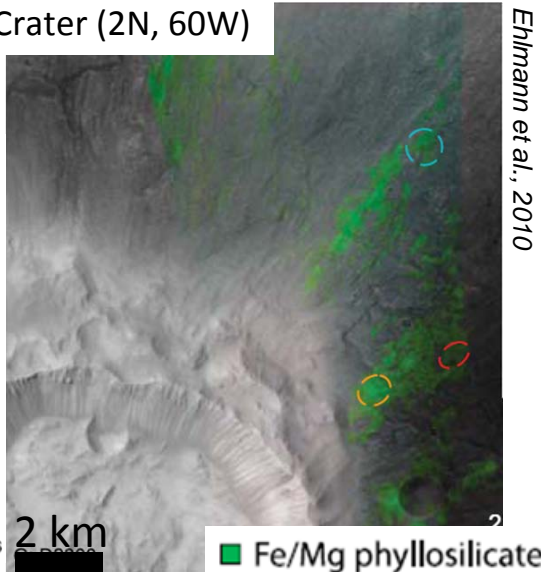
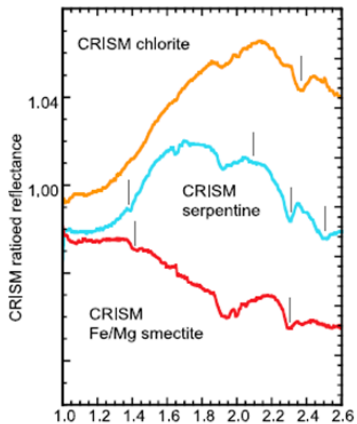
“Crustal Clays”

i. Exposed Ancient Crust, e.g. in the Claritas Rise



Ehlmann et al., 2010

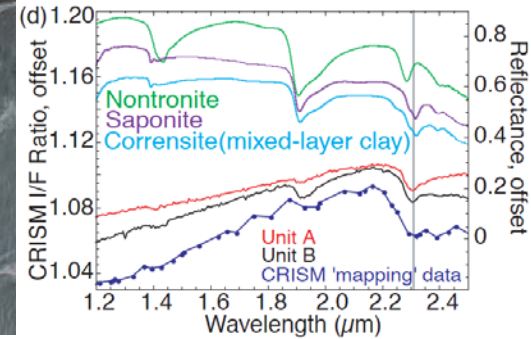
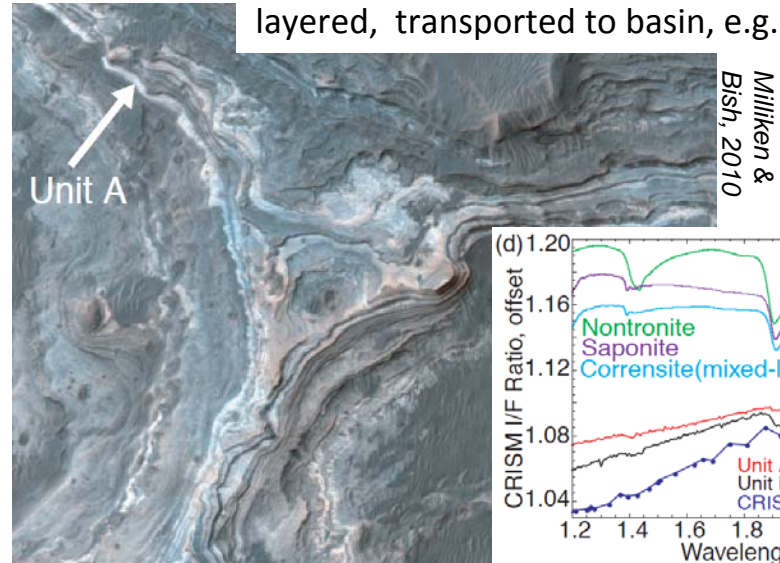
ii. Craters, e.g. Chia Crater (2N, 60W)



Ehlmann et al., 2010

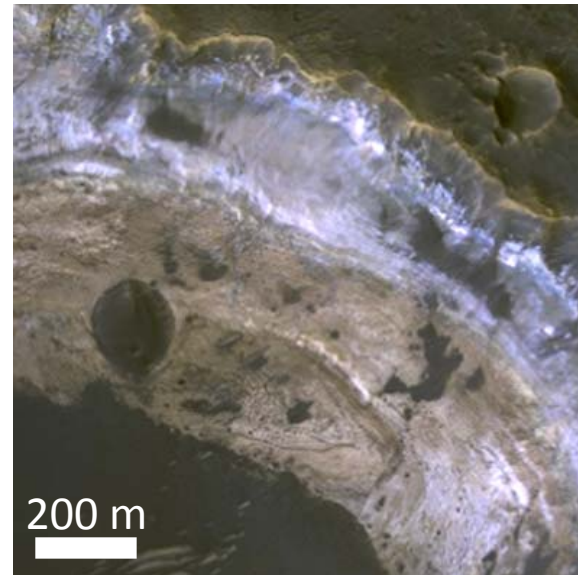
“Sedimentary Clays”

layered, transported to basin, e.g. Holden Crater



“Clays in Stratigraphy”

well-ordered compositional variation, e.g. Mawrth Vallis



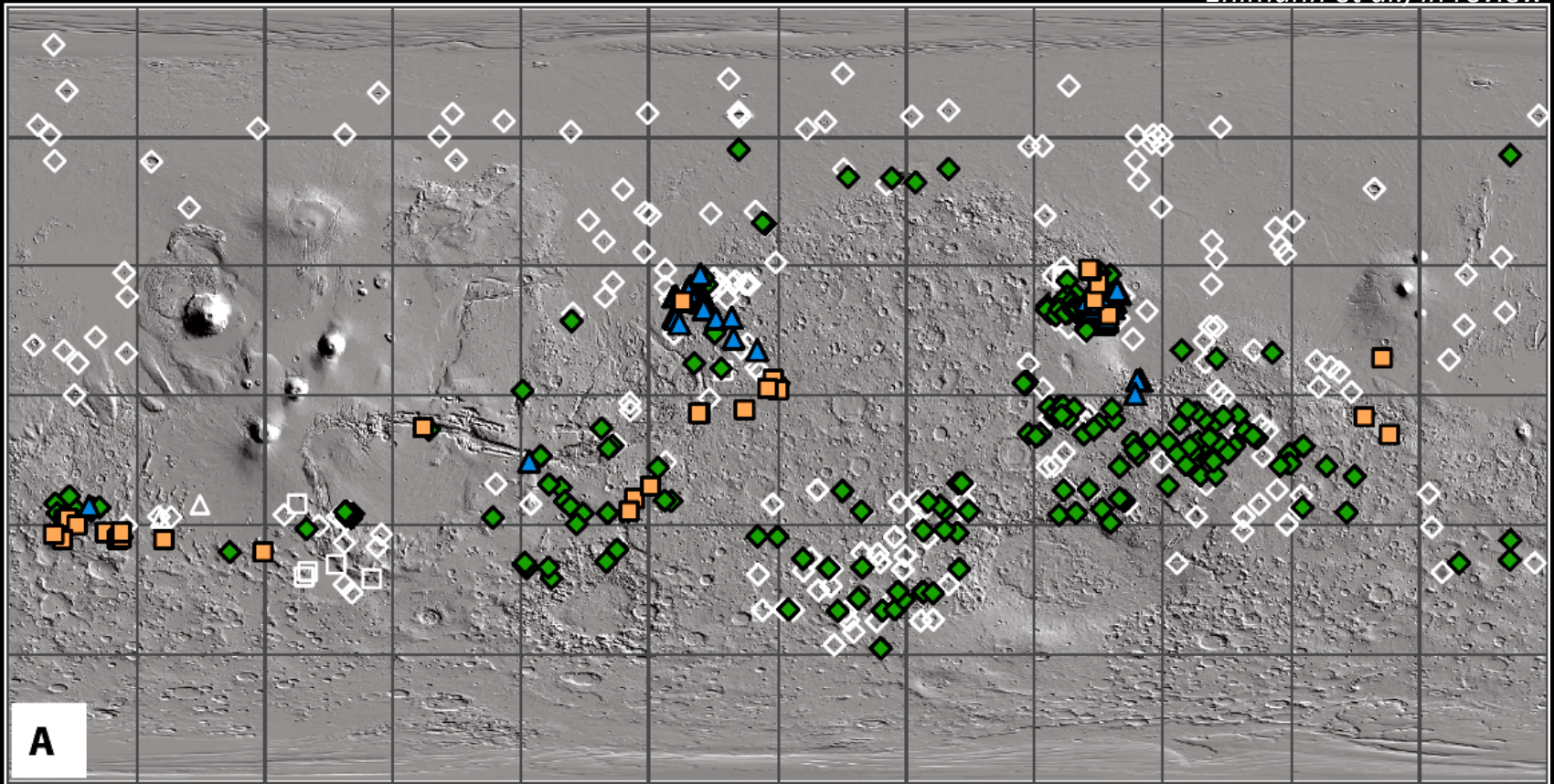
Loizeau et al. 2011

Al-clays +
silica

nontronite

Global Distribution of Clay-Bearing Materials

Ehlmann et al., in review



◆ = crustal clays

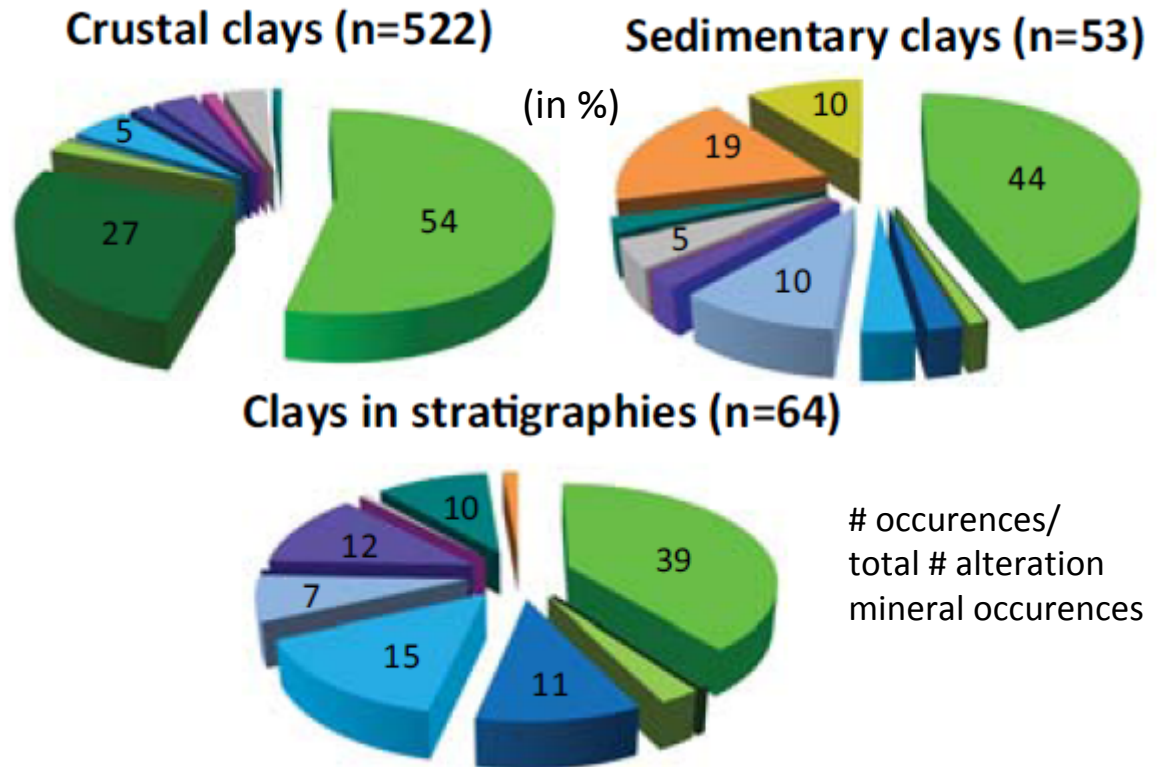
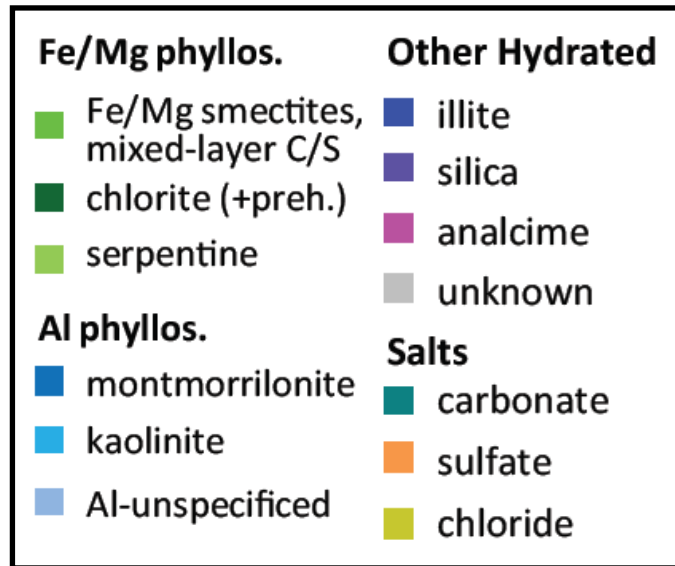
■ = sedimentary clays

▲ = clays in stratigraphy

open symbols = no clays found

→ Exposure is the principal control on mineralogy
(alteration to clays was a global phenomenon)

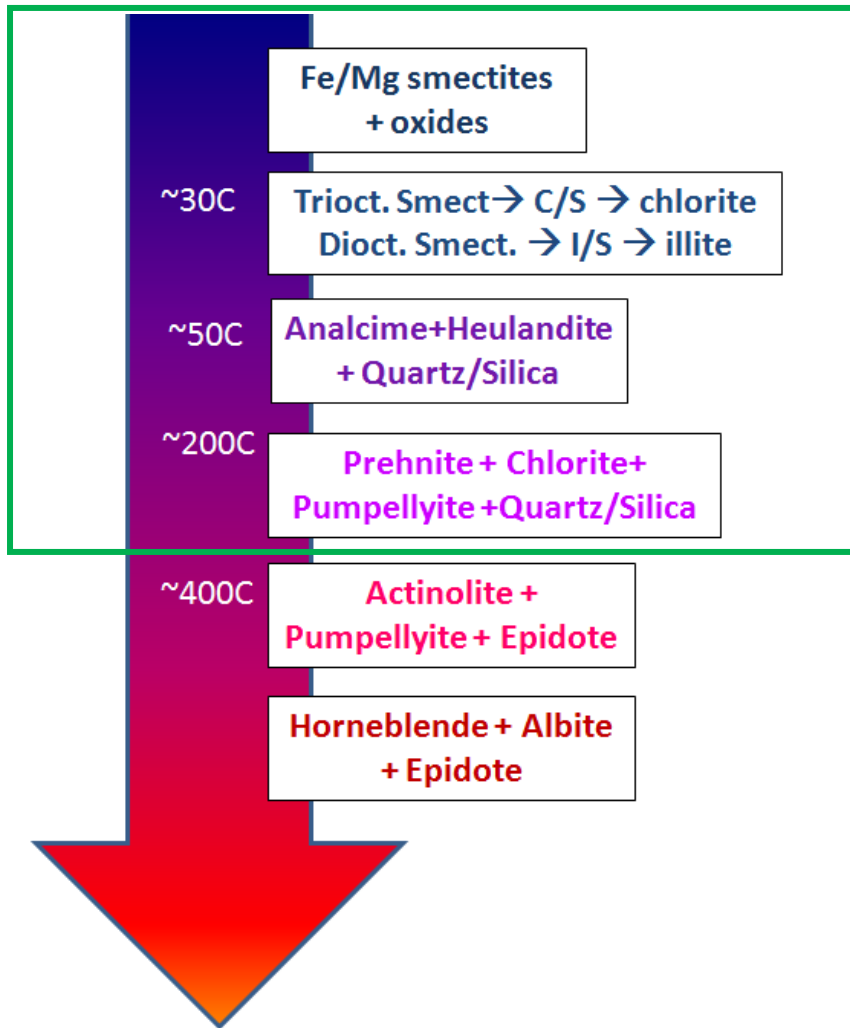
Mineral Associations: Key Findings



Ehlmann, et al., in review

- Fe/Mg smectites are always the most dominant clay mineral
- Phases associated with Fe/Mg smectites depend on geologic setting
 - Crustal: other hydrated silicates (prehnite, illite, analcime, silica, Al clays)
 - Sedimentary: salts + clays, more Al clays
 - in Stratigraphy: Fe/Mg smectites in stratigraphy with Al-clays or olivine-carbonate unit

LOW W:R ALTERATION OF BASALTIC PROTOLITH



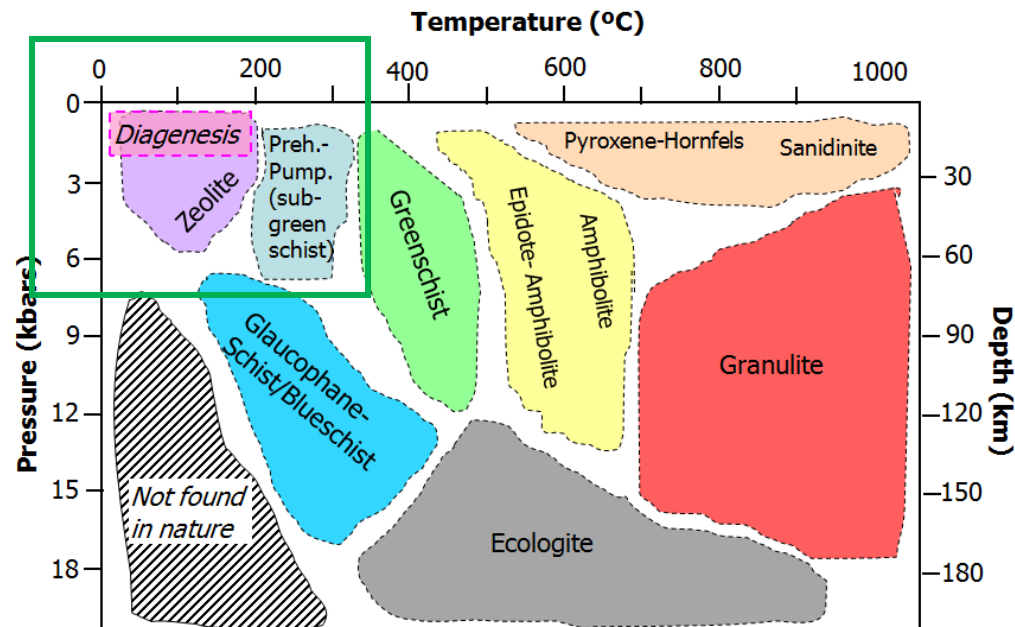
(an ultramafic composition assemblage would instead have saponite, serpentine, talc, brucite)

synthesizing Cann, 1979; Meunier, 2005; Frey & Robinson, 1999; Arkai et al., 2003; Philpotts & Aguee, 2009

Crustal Clays: Evidence of Elevated Temperature Alteration

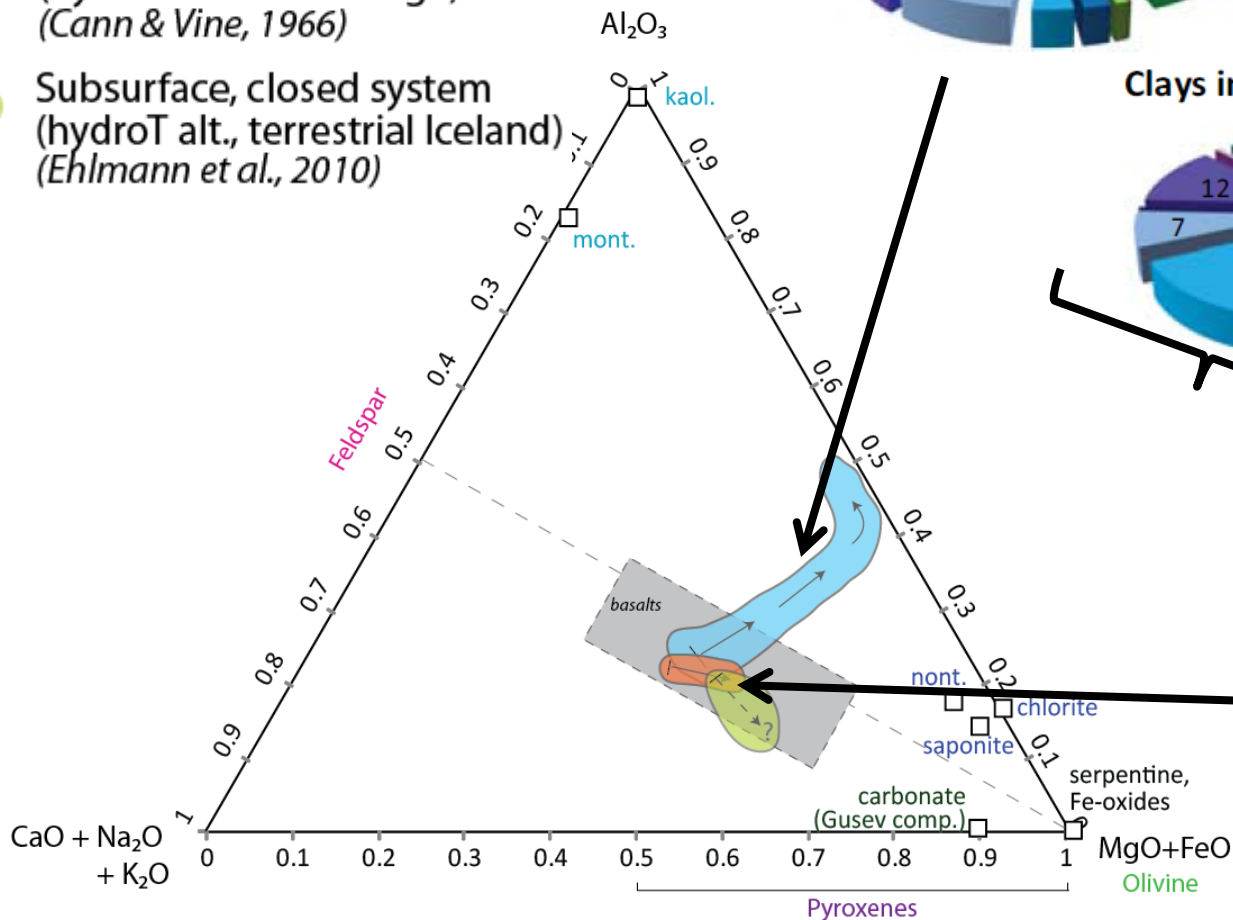
- Mineral assemblages are consistent with low W:R ratio formation by
 - Hydrothermal alteration
 - Low-grade metamorphism
 - Deuteric formation (e.g. lava cooling)

Ehlmann et al., forthcoming, Clays & Clay Minerals

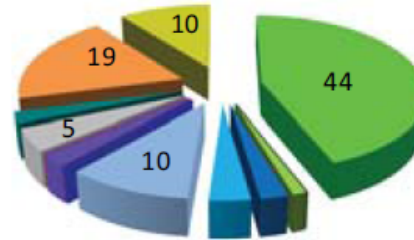


Styles of Alteration to form Clays

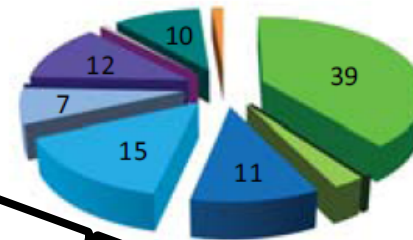
- Open system, high W:R, near surface
(Nesbitt & Wilson, 1992)
- Subsurface, closed system
(hydroT alt. ocean ridge)
(Cann & Vine, 1966)
- Subsurface, closed system
(hydroT alt., terrestrial Iceland)
(Ehlmann et al., 2010)



Sedimentary clays (n=53)



Clays in stratigraphies (n=64)

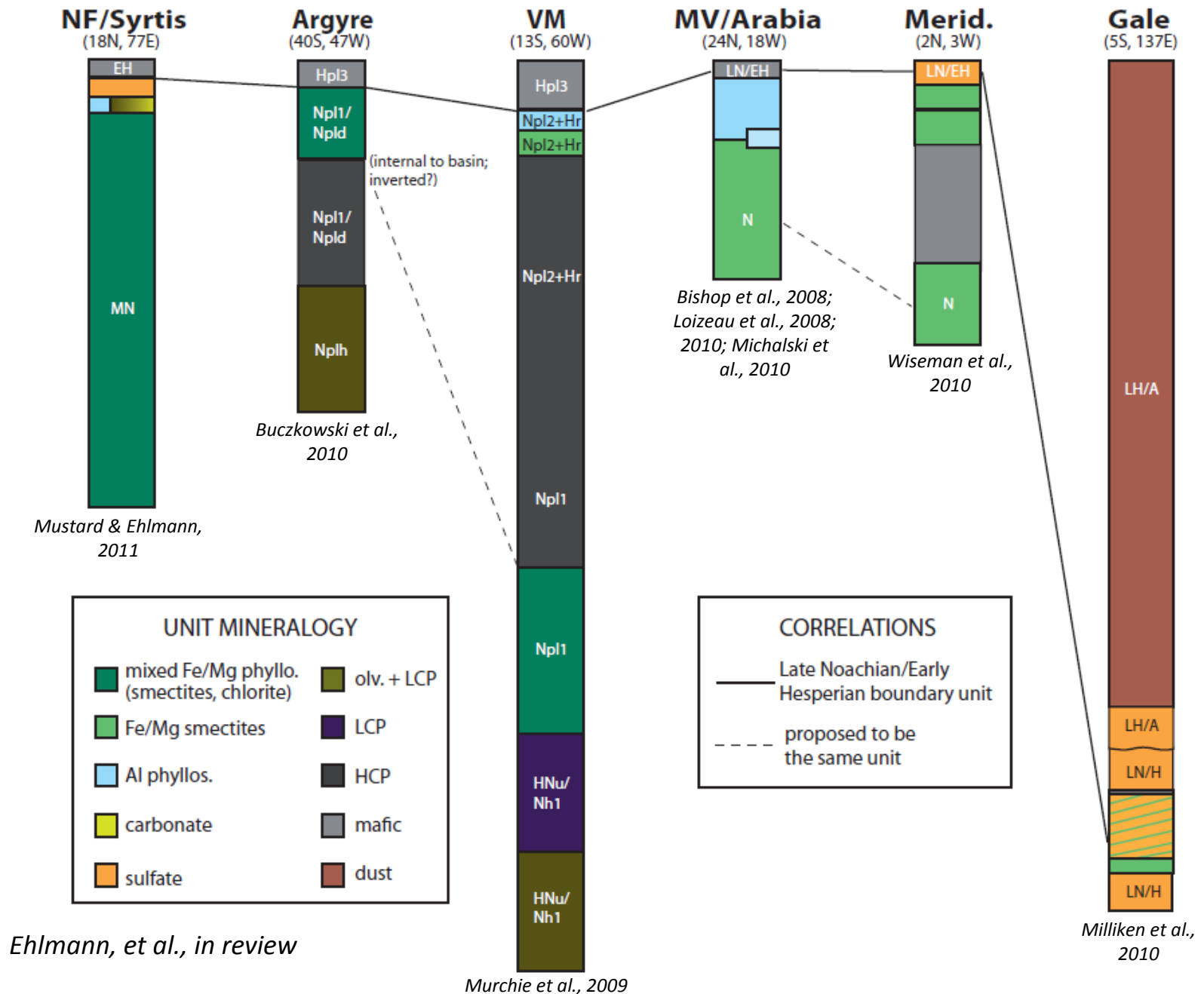


Open system alteration of crustal clays

Crustal clays (n=522)



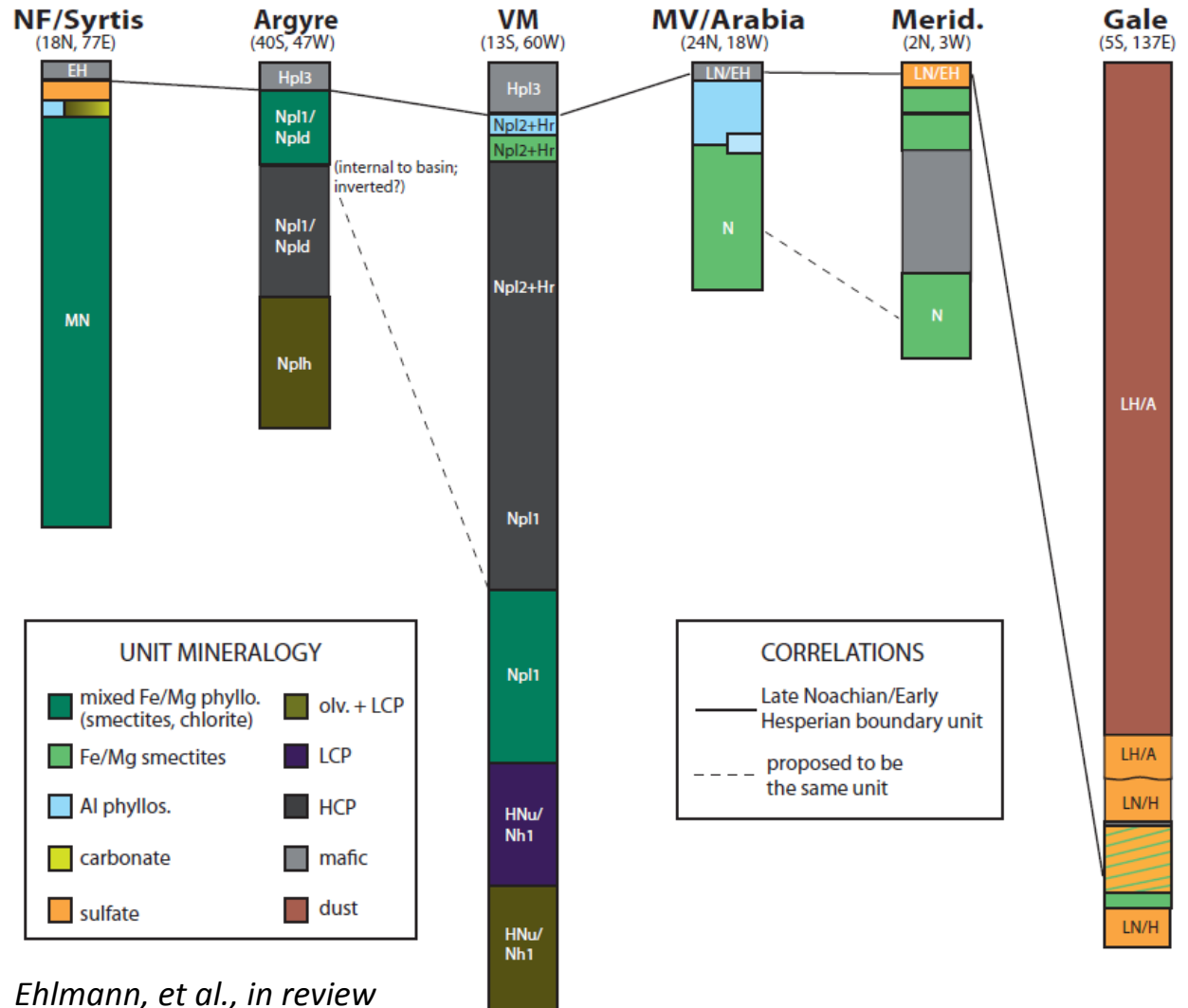
Compositional Stratigraphy of Clay-Bearing Units

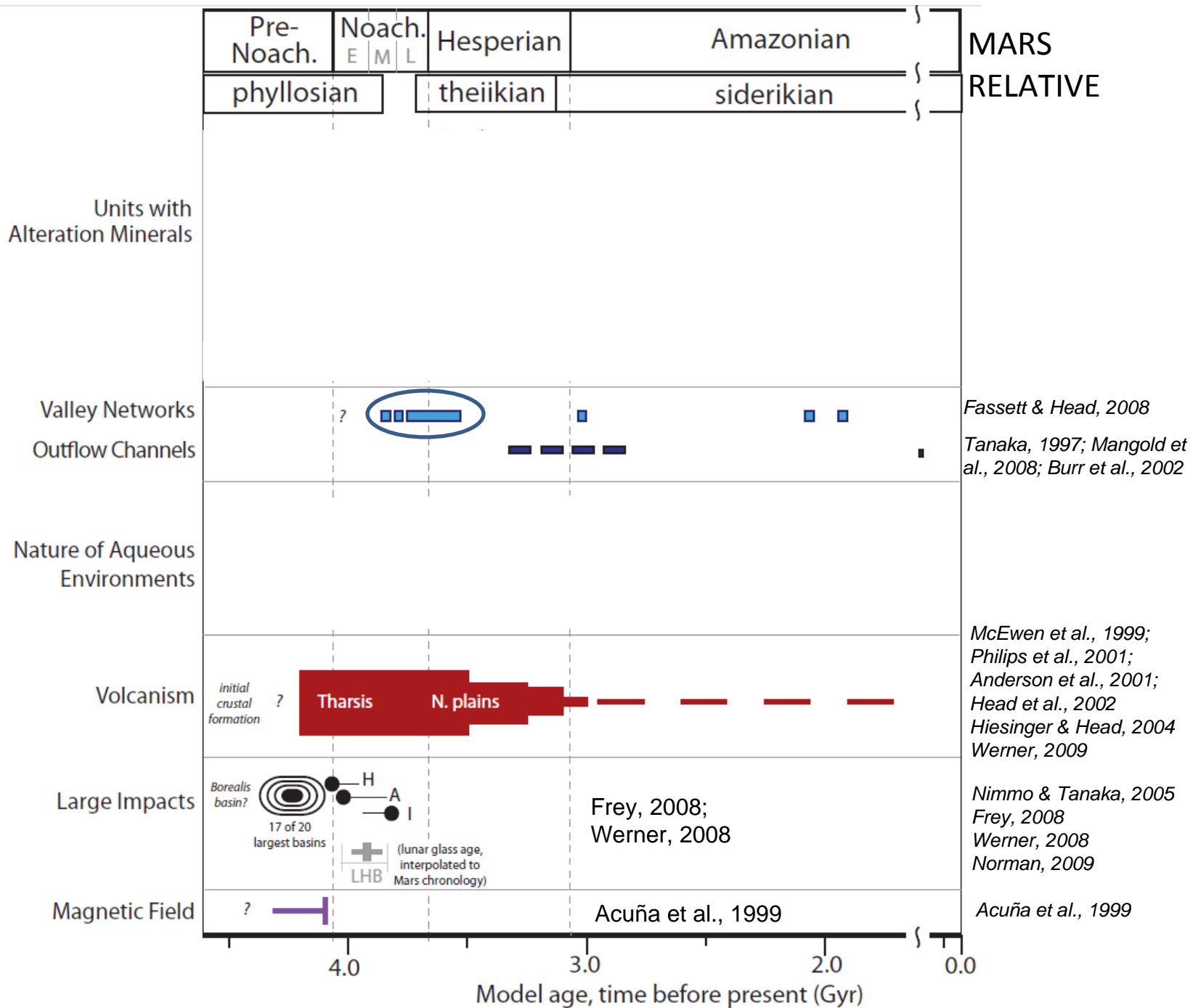


KEY POINTS

- No clays after crater count age corresponding to LN/EH
- Al clays always atop Fe/Mg phyllosilicates
- Fe/Mg phyllos. (smectite, chlorite, mixtures) comprise thick lower layers
- *but* not usually volumetrically abundant and many thick, deep units have no signs of alteration visible from orbit

Compositional Stratigraphy of Clay-Bearing Units





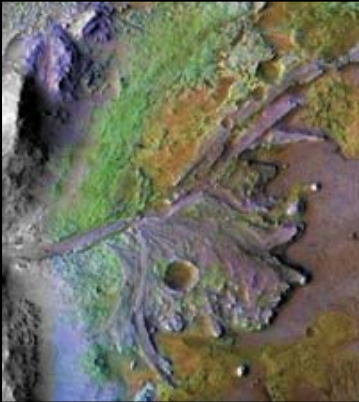
Mars Sedimentary Clays

Timing of all of these is predominately: Late-Noachian to Early-Hesperian
[Fassett & Head, 2005; 2008a; 2008b – Glotch et al., 2010 – Osterloo et al., 2010 – Wray et al., 2011 – Swayze, Ehlmann, Milliken, Poulet et al., in prep.]

Diversity in alteration products, settings implies diversity in water chemistry

JEZERO

Open-basin
Fe/Mg clay,
Mg carbonate



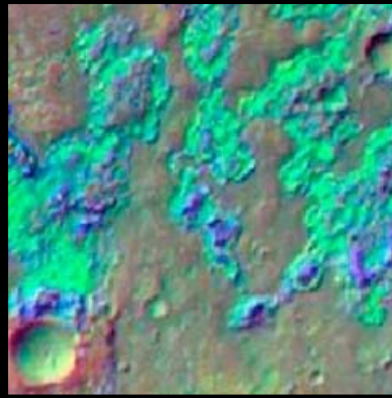
EBERSWALDE, HOLDEN

Open?-basin
Fe/Mg clay



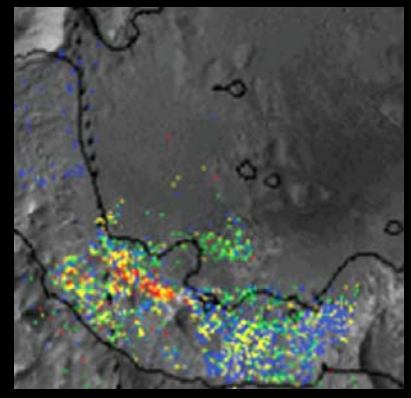
TERRA SIRENUM BASINS

Closed?-basin
Chloride salts
overlying, Fe/Mg
clays



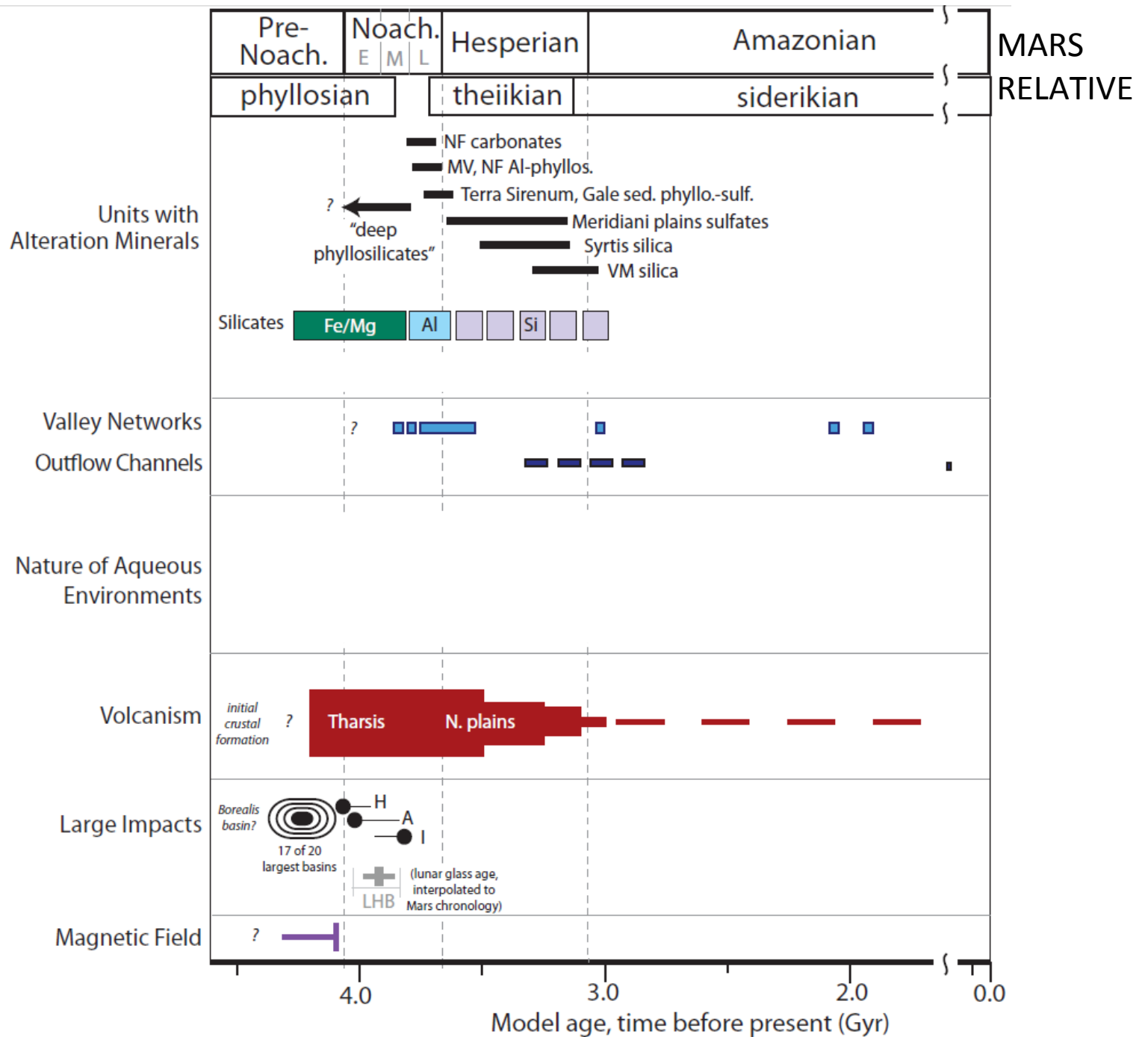
COLUMBUS , CROSS

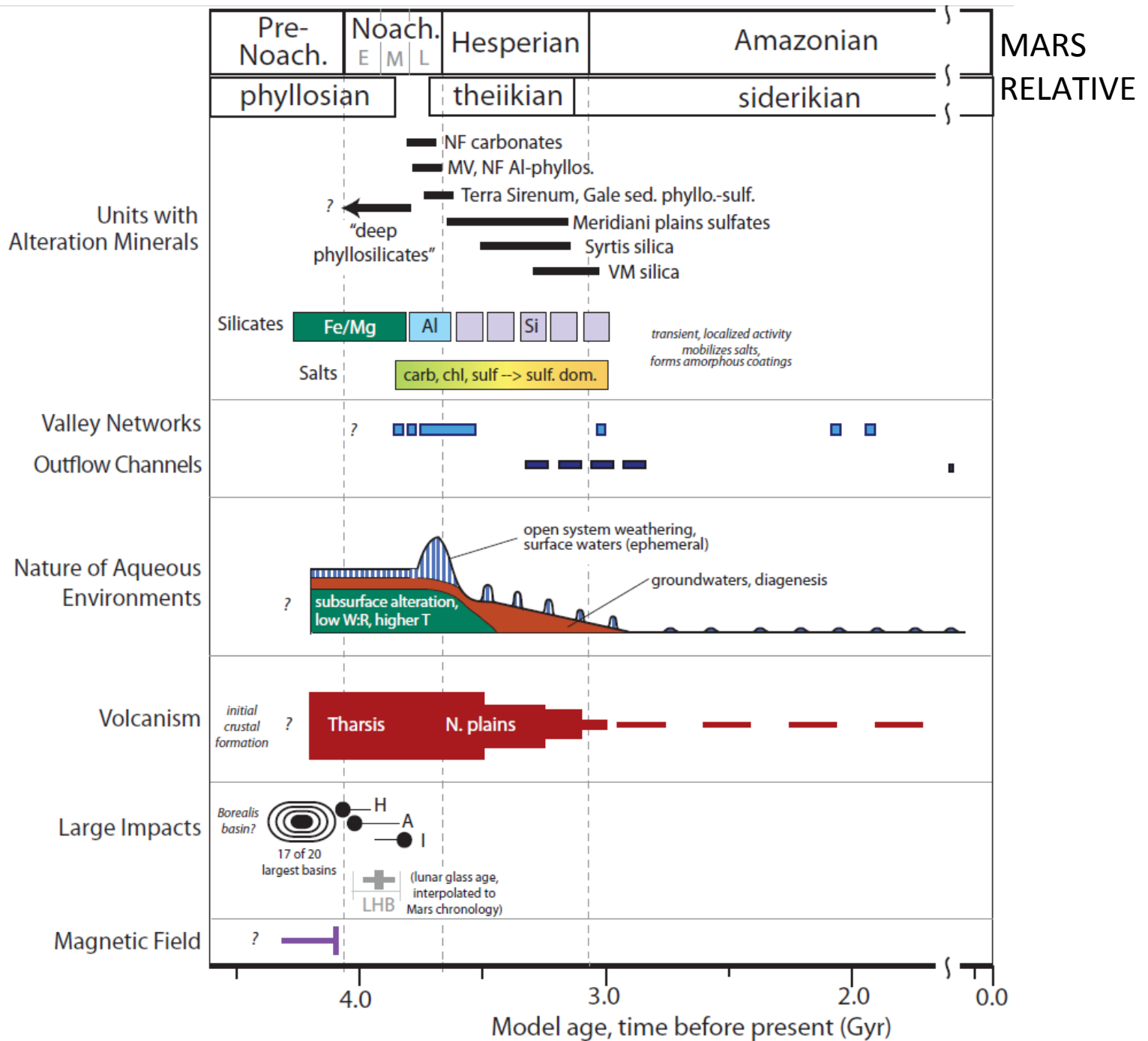
Closed-basin
Al (, Fe/Mg) clays
and sulfates (incl.
jarosite, alunite)



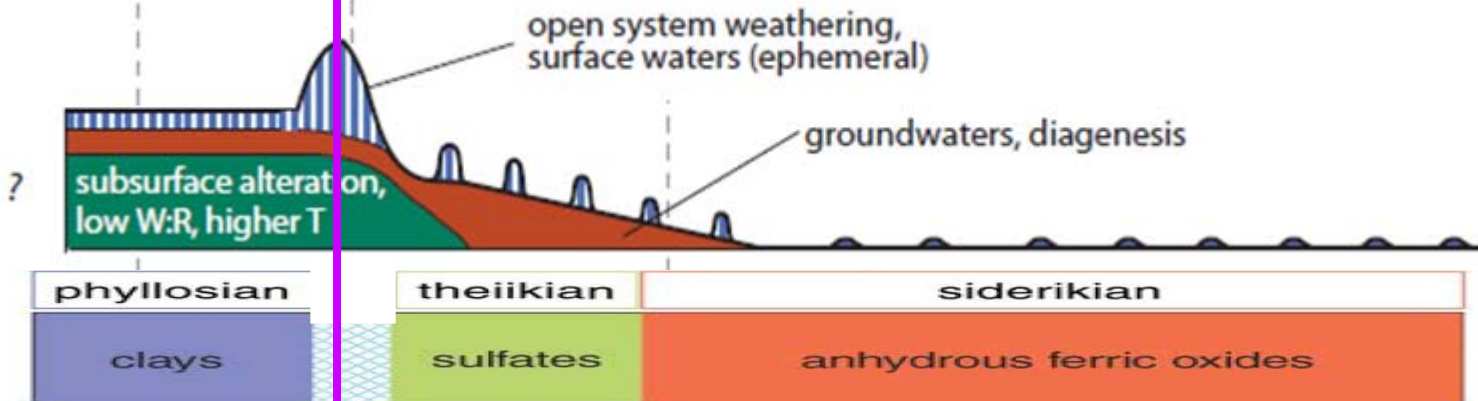
Alkaline

Acidic





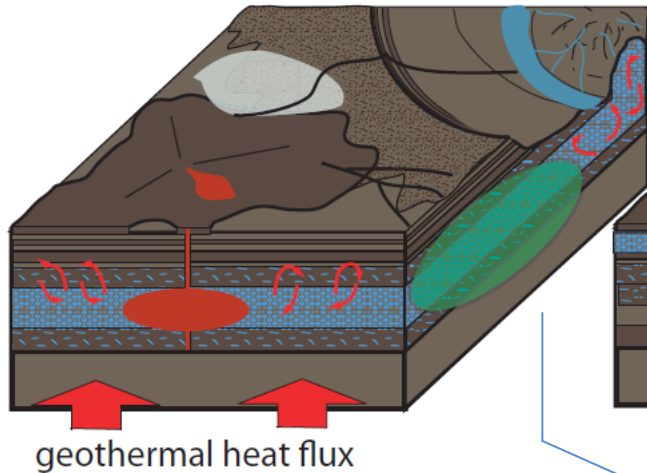
Noachian Hesperian Amazonian



- Most clay minerals—Fe/Mg phyllosilicate-bearing units deep in the crust—formed in the subsurface at low W:R in closed systems at temperatures ranging from ambient to low-grade hydrothermal (<300°C).
- Mechanism is consistent with
 - observed mineral assemblages of Fe/Mg smectites and chlorite, which form under anoxic, alkaline, high pH conditions
 - presence of the accompanying phases prehnite, analcime, serpentine and illite or muscovite, which form in diagenesis and/or in low-grade metamorphic/hydrothermal systems
 - the pervasive presence of clay minerals in large volumes of Mars' deepest exposed crustal materials (globally widespread) in spite of
 - elemental abundances and thermal infrared data consistent with dominance of primary minerals by weight fraction (*Taylor et al., 2010; Glotch et al., 2010*)
 - the morphology of Noachian fluvial valleys indicating hyperarid surface conditions except during the late Noachian, when episodic, high-magnitude floods incised numerous highland valley systems (*Irwin et al., 2005; Howard et al., 2005; Barnhart et al., 2009*).

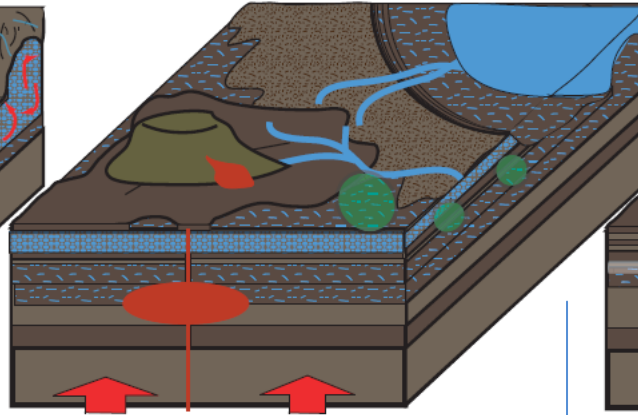
Inferred Environments

1. Global **phyllosilicate formation** (MN. & earlier)



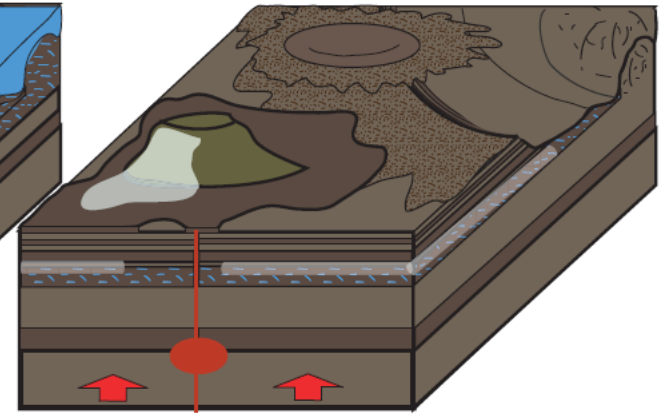
- Largely closed crustal system, low W:R ratio
- Assemblages of diverse phases with Fe/Mg phyllosilicates in low-grade hydrothermal/metamorphic/diagenetic conditions
- permits but does not require significant surface water**

2. Late-stage enhanced surface waters/alteration (LN/EH)



- High W:R open-system leaching (ice/permafrost melt, precipitation)
- Varying acidity
- Al clay + silica formation near the surface at large scales
- precipitation of various salts

3. Modern Mars (LH→present)



Ehlmann et al., in review

- colder, dryer
- very small amounts of water (in time or space)
- only local scale salt and silica precipitation/mobility

Further Hypothesis Testing: 'Most Clays Formed in the Subsurface'

1. **direct *in-situ* detection of hydrothermal/metamorphic indicator minerals or interstratified clays**
2. **elemental abundances of clay-bearing materials are little-changed from that of basalt (indicative of low W:R style)**
3. **textures of clay-bearing materials with alteration phases concentrated in veins and vesicles (similar to that observed in Mars meteorites)**
4. **isotopic data (from H, C, O) show elevated formation temperatures for minerals**
5. (indirect) measured isotopic ratios and calculation of loss rates of **atmospheric gases that indicate a continuously thin atmosphere**

Implications

- Early enhanced rates of crustal cooling
 - Explanation for early topographic feature preservation, e.g. Parmentier & Zuber, 2007
- Possibly significant sequestration of volatiles in the crust
- **Early atmosphere does not need to be thick, nor surface warm**
 - Transient liquid waters only, (e.g. from large-scale volcanic-gas release: Johnson et al., 2008; Phillips et al., 2001)
- **Crustal environments may be the oldest, most stable, and longest aqueous (habitable) environments on Mars**
 - Habitability of low W:R hydrothermal/metamorphic systems?
 - Concentration of the products of life for detectability?
 - Mars' biosphere is/was a deep biosphere?

EXTRAS

For CRISM images with clay-bearing materials, % with...

	<i>Alteration Phase</i>	Crustal (% sites)	Sedimentary (% sites)	in Stratigraphy (% sites)
Fe/Mg CLAYS	<i>Fe/Mg smectite</i>	77.6	84.4	95.1
	<i>Chlorite/prehnite</i>	39.0	0	1.6
	<i>Serpentine</i>	3.1	0	6.6
Al CLAYS	<i>Montmorillonite</i>	3.9	4.4	26.2
	<i>Kaolinite</i>	7.7	6.7	36.1
	<i>Al-phylo (unspecified)</i>	0	20.0	18.0
OTHER HYDRATED SILICATE	<i>Illite</i>	2.3	0	0
	<i>Silica</i>	5.4	4.4	29.5
	<i>Analcime</i>	1.9	0	0
	<i>Other Hydrated (unspecified)</i>	5.4	8.9	1.6
SALTS	<i>Carbonate</i>	1.2	4.4	24.6
	<i>Sulfate</i>	0	35.6	3.3
	<i>Chloride</i>	0	20.0	0

Mineral Indicators of Elevated Temperature

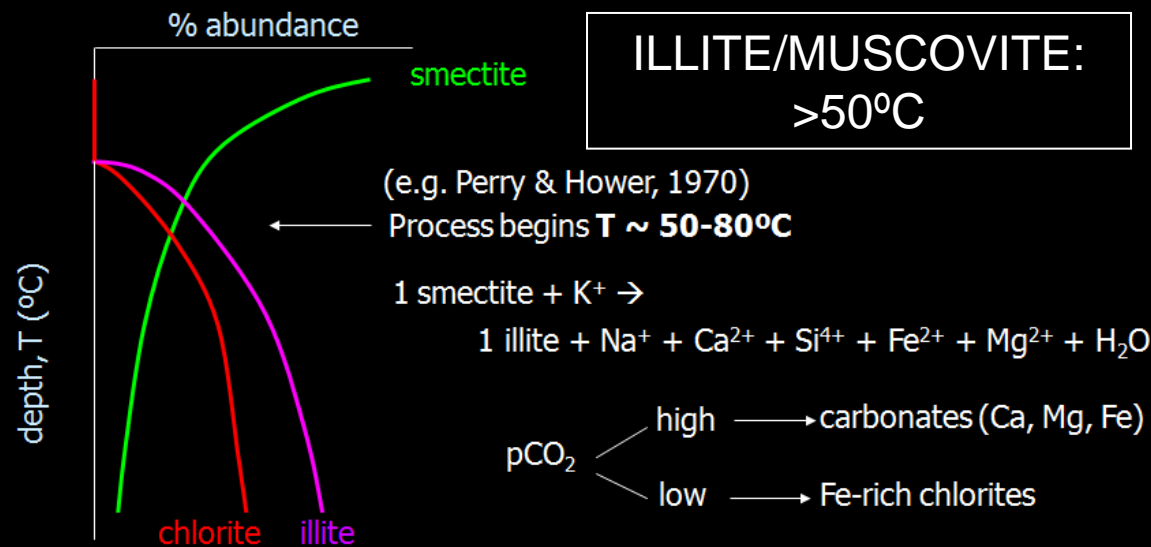
ANALCIME: 50-250°C



Zeolite zone	Chabasite-Thomsonite Zone	Mesolite-Scolecite Zone	© Tobias Weisenberger Laumontite Zone		
Mineral	50°C	100°C	150°C	200°C	250°C
Levyn	50-100				
Chabasite	50-100				
Thomsonite	50-100				
Gismondine	50-100				
Mesolite-Scolecite		100-150			
Heulandite		100-150			
Phillipsite		100-150			
Stilbite		100-150			
Epistilbite		100-150			
Mordenite		100-150			
Laumontite		100-150			
Analcime		100-150			

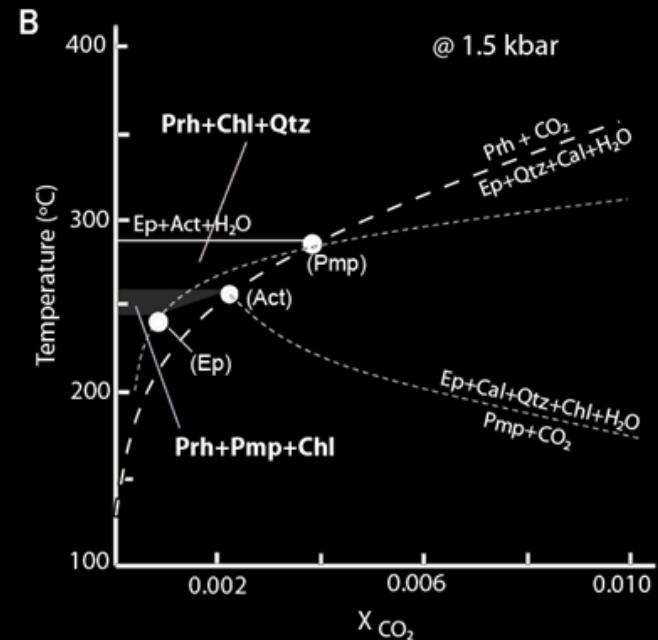
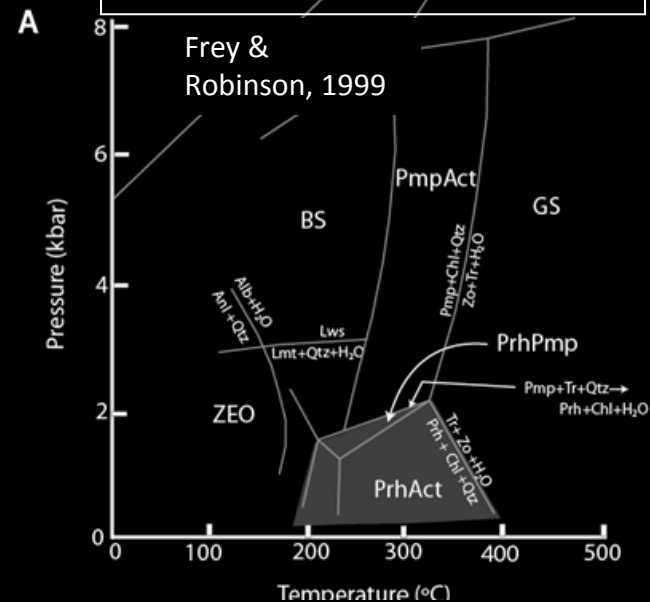
(Weisenberger & Selbekk, 2008)

Zeolitized basalt, Iceland



Ehlmann et al., JGR, 2009; CCM, in review

PREHNITE: 200-400°C



VNIR Quantitative Mineral Modeling

Little evidence of chemical fractionation in clay-bearing units

